

锂同位素示踪大陆风化：进展与挑战

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摘要: 大陆风化制约着地表物质循环及其从陆地向湖泊 / 海洋的迁移，并通过消耗大气 CO₂ 调节长时间尺度的全球碳循环和气候变化，因此如何有效示踪大陆风化是地球表生过程研究的重要科学问题之一。锂 (Li) 的两个同位素 (⁶Li 和 ⁷Li) 拥有巨大相对质量差、无化合价变化，且不受氧化还原条件和生物作用影响等优势，赋予 Li 同位素体系具备示踪大陆风化的潜力。然而，风化体系中 Li 的来源和分馏的制约要素争议颇多。本文从储库、风化壳、河流体系、淋滤实验和模型模拟等方面综述了目前 Li 同位素示踪大陆风化的研究现状和存在的挑战。最后指出，在示踪大陆风化方面，Li 同位素具有其独特的作用，建议细化岩石 / 矿物淋滤实验、大小流域相结合、加强多同位素体系相互补充与验证、以及计算机模型模拟，有望减少 Li 同位素示踪大陆风化的不确定性。

关键词: 锂同位素；大陆风化；地球化学示踪；制约要素；综述

Using lithium isotopes traces continental weathering: Progresses and challenges

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Abstract: Background, aim, and scope Since weathering processes constrain the cycle of elements and their transportation from the continents to the lake/ocean reservoirs, shape the topography of continents, and regulate global carbon cycle and therefore climate changes by consuming atmospheric carbon dioxide over geological time-scales, how to effectively trace continental weathering processes is one of key scientific topics of supergene geochemistry. Being one of the most promising silicate-weathering tracers, lithium (Li) isotopes have long been exploited to trace continental weathering and thus to reconstruct secular weathering scenarios. To acquire effective research methodology, review of establishing Li isotopes to be a silicate-weathering tracer is thus essential. **Materials and methods** This paper reviewed current major developments, problems and challenges of using Li isotopes to trace weathering with respects to its major reservoirs, weathered crust, riverine systems,

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leaching experiments, and modeling simulation which have progressively developed in the past decades. **Results** Owing to large mass difference between ${}^6\text{Li}$ and ${}^7\text{Li}$, one valence, and no effect of redox conditions and living beings, if not all, at least at catchment scale, Li isotopes own potential to trace continental weathering processes. However, its sources and controlling factors of Li fractionation during weathering processes are controversial. **Discussion** Changes in Li isotopic compositions from weathered crust to fresh bedrocks are complex and even reverse, whereas controlling factors for Li isotopic variations in riverine system are various and controversial, in particular for large river systems. Leaching experiments and modeling simulation for Li isotopic compositions are at an early stage. More attentions are focused on seasonal variation in riverine Li isotopic compositions at small, monolithological catchment. **Conclusions** We pointed out that Li isotope system does have its potential and uniqueness in tracing continental weathering, especially for silicate weathering. **Recommendations and perspectives** In order to reduce the uncertainty of tracing silicate weathering processes using Li isotopes, elaborating leaching experiments of rocks/minerals, marrying large and small catchments, enhancing computation simulation, and coupling of multi-isotopes are suggested to be needed in future.

Key words: Li isotopes; continental weathering; geochemical tracing; controlling factors; review

锂(Li)同位素的研究起始于 McLennan and Ainslie(1922), 继 Taylor and Urey(1937, 1938)讨论了Li同位素分馏之后, 由于分析测试方法的限制, Li同位素的研究一直发展缓慢。直至20世纪80年代以来, 得益于质谱技术的革新, 如热电离质谱仪(TIMS)、二次离子质谱仪(SIMS)、离子探针(Ion Probe), 尤其是多接收电感耦合等离子体质谱仪(MC-ICP-MS)的出现(赵葵东和蒋少涌, 2001; Su et al., 2015), Li同位素的测试精度已达到0.2‰左右(Huang et al., 2010; Van Hoecke et al., 2015; Lin et al., 2016), 其SIMS微区原位分析精度也达到1‰左右(李献华等, 2015), 致使Li同位素的研究迅速发展。目前, Li同位素已被广泛应用于示踪星云、行星的起源和演化(Seitz et al., 2004, 2007; Liu et al., 2009)、壳-幔物质循环(Tomasca et al., 2000; Magna et al., 2002; Su et al., 2014; Zheng et al., 2015)、板块俯冲及岛弧演化(Zack et al., 2003; Tian et al., 2015; 万红琼等, 2015)、盐湖形成与演化(肖应凯等, 1994; Orberger et al., 2015)、成矿物质来源(Helvaci et al., 2004; 苏媛娜等, 2011; Dill and Weber, 2013)、重建地质尺度的环境变化(Froelich and Misra, 2014; Li and West, 2014; Vigier et al., 2015)以及行星和大陆风化(Ushikubo et al., 2008; Liu and Rudnick, 2011; Wang et al., 2015; Fairén et al., 2015)等。在示踪大陆风化作用方面, Li同位素为化学风化过程和物质循环研究注入了全新的活力(Vigier et al., 2009;

Misra and Froelich, 2012; Dellinger et al., 2015)。

由于硅酸盐岩的化学风化被认为是地质时间尺度上最重要的碳汇过程(Berner et al., 1983; Kump et al., 2000), 而Li同位素具有揭示硅酸盐岩风化强度和通量、示踪地壳物质循环(特别是营养盐循环)的潜力(Berner and Kothavala, 2001; Vigier et al., 2009; Misra and Froelich, 2012; Cermeno et al., 2015), 因此Li同位素被用来示踪岩石/矿物化学风化、全球/区域环境变化、构造运动、生物演替等过程, 进而为理解大陆风化及其对构造和气候演化的响应和反馈机制提供了全新的视角。然而, 随着对Li同位素示踪大陆风化研究的拓展和深入, 流域Li同位素分馏控制要素颇具争议, 风化壳Li同位素难以类比等问题愈加突出。为了更好地理解Li同位素在示踪大陆风化方面的优势及局限, 本文综述了Li同位素在风化壳、流域风化、岩石实验淋滤、模型模拟等方面的研究进展及其面临的挑战, 以期促进Li同位素在示踪大陆风化方面的进一步发展。

1 大陆岩石和风化产物Li含量及其同位素组成

1.1 自然储库中的Li同位素组成

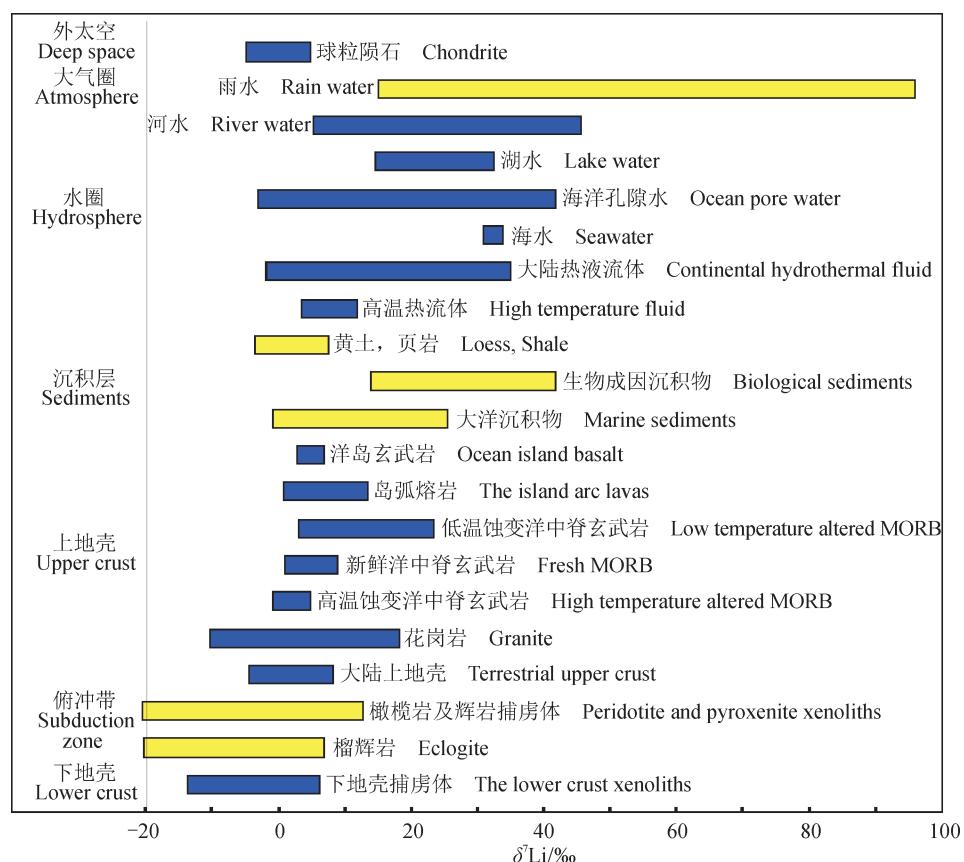
原子序数为3的Li有 ${}^6\text{Li}$ 和 ${}^7\text{Li}$ 两个质量差约17%的稳定同位素, 其同位素组成的表达式为:

$$\delta^7\text{Li}(\%) = \frac{\frac{7}{6}\text{Li}(\text{样品}) - \frac{7}{6}\text{Li}(\text{LSVEC})}{\frac{7}{6}\text{Li}(\text{LSVEC})} \times 1000\% \quad (1)$$

目前广泛使用的 Li 的标准物质是美国国家标准与技术研究所合成的碳酸锂 (Li_2CO_3) , 即 NIST-LSVEC, 其 ${}^7\text{Li}/{}^6\text{Li}=12.1025\pm 0.0016$ (Chan et al, 2009)。

Li 具有中等不相容的性质, Li^+ 半径为 0.68 \AA , 与 Mg^{2+} (0.66 \AA)、 Al^{3+} 半径 (0.51 \AA) 和 Fe^{2+} (0.74 \AA) 的半径比较接近, 因此 Li^+ 可以部分替代铁镁质岩石中的 Mg、Fe (刘英俊等, 1984)。从基性岩到酸性岩, Li 含量升高, 例如辉长岩中 Li 含量是 $1.5 \mu\text{g}\cdot\text{g}^{-1}$, 橄榄岩中 Li 的含量是 $4 \mu\text{g}\cdot\text{g}^{-1}$ (Vils et al, 2008), 花岗岩的 Li 含量在 $6\text{--}40 \mu\text{g}\cdot\text{g}^{-1}$ (刘英俊等, 1984)。不同岩性的 $\delta^7\text{Li}$ 值变化也较明显, 辉长岩中 Li 同位素约为 4\textperthousand , 橄榄岩中 Li 同位素组成约为 4\textperthousand (Vils et al, 2008), 花岗岩的 Li 同位素组成约为 $-10\text{\textperthousand}$ — $18\text{\textperthousand}$ (Tomascak, 2004)。

Li 同位素体系具备的特有示踪潜力使其发展迅速, 目前已经基本获得了自然储库中 Li 含量及 $\delta^7\text{Li}$ 值 (图 1, 表 1), 这为 Li 同位素示踪大陆风化的研究奠定了基础 (Tomascak, 2004)。大洋中脊新鲜玄武岩 (MORB) 的 Li 含量及其同位素组成分别是 $3.1\text{--}7.5 \mu\text{g}\cdot\text{g}^{-1}$ 和 $1.5\text{\textperthousand}$ — $5.6\text{\textperthousand}$ (Chan et al, 1992; Liu et al, 2011)。在岩浆结晶过程中, Li 优先进入流体相 (Chan et al, 1994; You and Chan, 1996; Brenan et al, 1998), 导致其在地幔亏损, 而在地壳富集。大陆地壳主要成分是花岗闪长质, 在近地表处接近花岗岩成分 (Mason and Moore, 1982; Wedepohl, 1995; 韩玲文等, 2003), 由此上地壳中 Li 的平均含量及其同位素比值接近花岗岩, 含量约为 $30.5 \mu\text{g}\cdot\text{g}^{-1}$, $\delta^7\text{Li}$ 约为 $0.6\text{\textperthousand}$ (Tomascak, 2004; Misra and Froelich, 2012; Sauzéat et al, 2015)。



补充数据: 河水引自文献 Henchiri et al (2014) 和 Wang et al (2015), 雨水则汇总了文献肖应凯等 (1994)、Pogge von Strandmann et al (2006)、Vigier et al (2009) 和 Millot et al (2010a, 2010b, 2010c) 等文献。

Supplementary data of river waters are from Henchiri et al (2014) and Wang et al (2015) and that of rain waters from Xiao et al (1994), Pogge von Strandmann et al (2006), Vigier et al (2009) and Millot et al (2010a, 2010b, 2010c).

图 1 自然储库中 Li 同位素的分布 (据汤艳杰等 (2009) 补充)

Fig.1 Lithium isotopic compositions in natural reservoirs (Modified from Tang et al (2009))

表1 大陆常见物质及相关风化产物的 Li 含量及其同位素组成
Tab.1 Lithium concentrations and isotopic compositions of fresh and weathered crust materials on the continents

物质类型 Type	名称 Name	含量 Concentration/($\mu\text{g}\cdot\text{g}^{-1}$)	$\delta^7\text{Li}/\text{\textperthousand}$	数据来源 Sources of data
常见上地壳物质 Common upper crust materials	黄土 Loess	~30.5	~0.6	Tsai et al., 2014; Sauzéat et al., 2015
	土壤 Soil	5—200	-2.6—14	Ryu et al., 2014
	花岗岩 Granite	6—40	-10—18	Tomascak, 2004
	新鲜 MORB Fresh MORB	3.1—7.5	1.5—5.6	Chan et al., 1992; Liu et al., 2011
	橄榄岩 Peridotite	~4	~4	Vils et al., 2008
风化相关产物 Related weathered products	粘土矿物 Clays	~80	-1.6—5	
	陆源碎屑 Terrigenous clast	~24	0—6	Hoefs and Sywall, 1997;
	生物成因沉积物 Biogenic sediments	很低 Very low	6—32	Chan and Hein, 2007; Vils et al., 2008;
	沉积碳酸盐岩 Sedimentary carbonates	0.2—4	-41—25	
	盐湖 Salt lakes	0—262		Yu et al., 2013

1.2 风化产物的 Li 含量及同位素组成

由于 Li 是水溶性元素, 因此 Li 将随风化淋溶作用迁移至溶液中。然而, 大量研究表明在风化淋溶过程中 Li 仅仅发生微弱的同位素分馏, 或者不发生显著的同位素分馏, Li 同位素的分馏主要发生在其搬运过程中 (Pistiner and Henderson, 2003; Qiu et al., 2009; Wimpenny et al., 2010a; Verney-Carron et al., 2011)。在搬运过程中, ${}^6\text{Li}$ 优先在次生粘土矿物中富集, 因此粘土物质中 Li 含量较高, 平均约为 $80 \mu\text{g}\cdot\text{g}^{-1}$, $\delta^7\text{Li}$ 值为 $-1.6\text{\textperthousand} - 5\text{\textperthousand}$; ${}^7\text{Li}$ 则主要随流体带入海洋, 这些 ${}^7\text{Li}$ 随俯冲过程被带到下地壳 / 地幔, 最后造成上地壳富 ${}^6\text{Li}$ (Marschall et al., 2007a, 2007b)。陆源碎屑沉积物中 Li 平均含量为 $24 \mu\text{g}\cdot\text{g}^{-1}$, $\delta^7\text{Li}$ 值则变化于 $0\text{\textperthousand} - 6\text{\textperthousand}$; 生物成因沉积物 Li 含量极低, 其同位素组成约为 $6\text{\textperthousand} - 32\text{\textperthousand}$; 沉积碳酸盐岩中 Li 含量一般小于 $5 \mu\text{g}\cdot\text{g}^{-1}$, 其 $\delta^7\text{Li}$ 值为 $-1.6\text{\textperthousand} - 5\text{\textperthousand}$ (Hoefs and Sywall, 1997; Chan and Hein, 2007; Vils et al., 2008)。新生代碳酸盐岩中的 Li 含量很低, 约在 $0.2 - 4 \mu\text{g}\cdot\text{g}^{-1}$, $\delta^7\text{Li}$ 值约为 $-41\text{\textperthousand} - 25\text{\textperthousand}$ (Hoefs and Sywall, 1997)。

盐湖中 Li 含量变化范围则非常大, 变化于 $0 - 262 \mu\text{g}\cdot\text{g}^{-1}$ (表1), 这与盐湖各自成因及不同的物质来源有关 (Yu et al., 2013) (表1)。

土壤中 Li 含量变化范围也比较大, 为 $5 - 200 \mu\text{g}\cdot\text{g}^{-1}$ (刘英俊等, 1984), $\delta^7\text{Li}$ 值在 $-2.6\text{\textperthousand} - 14\text{\textperthousand}$ (Ryu et al., 2014)。页岩中 Li 含量平均为 $60 \mu\text{g}\cdot\text{g}^{-1}$, 一些富 Li 矿物中 Li 含量可以高达 1% (Meier, 1982)。黄土中 Li 平均含量与上地壳基本一致, 变化于 $20 - 35 \mu\text{g}\cdot\text{g}^{-1}$, 平均值约为 $30.5 \mu\text{g}\cdot\text{g}^{-1}$, $\delta^7\text{Li}$ 值也较为均一, 约为 $0.6\text{\textperthousand}$ (Jia et al., 2007; Tsai et al., 2014; Sauzéat et al., 2015)。

2 Li 同位素示踪大陆风化研究进展

自然界中 Li 一般以 Li^+ 离子形式存在 (Faure and Mensing, 2005), 无化合价变化, 氧化还原条件对其分馏没有影响。同时 Li 不是生命元素, 生物作用对其分馏也可能没有影响 (Rudnick et al., 2004; Lemarchand et al., 2010; Clergue et al., 2015)。所有地质过程中 Li 同位素的最大分馏发生在大陆风化作用过程中 (Rudnick et al., 2004;

Liu and Rudnick, 2011; Henchiri et al, 2014) (图1), 在自然界已观察到超过 110‰ 的分馏, 约 90‰ 的 Li 同位素分馏发生在地表 / 浅地表环境 (Hoefs and Sywall, 1997; Rudnick et al, 2004; Yoon, 2009; Millot et al, 2010a)。与 (浅) 地表环境下约 90‰ 的 Li 同位素分馏相比, 直接由温度引起的 Li 同位素分馏则微不足道, 仅为 2‰ (Millot et al, 2010b; Wimpenny et al, 2010b)。因此, 拥有较大相对质量差的⁶Li 和⁷Li 具备示踪大陆风化作用的潜力 (Chan et al, 1992; Kisakürek et al, 2004a; Misra and Froelich, 2012; Dellinger et al, 2015)。

然而, 到底哪些因素影响着由化学风化导致的不同气候带河水 Li 含量及⁷Li 值变化呢? 此问题一直争议不断 (Huh et al, 1998, 2001)。为了回答这一问题, 早期研究者尝试对风化壳剖面进行剖析, 试图为 Li 同位素示踪大陆风化提供线索。然而, 由于风化壳自身成因的复杂性, 结果不太理想 (见下讨论)。随后, 为了认识 Li 同位素在大陆风化过程中的行为, 研究者开始对全球主要大河流和不同气候区小流域进行调查分析, 并开展室内淋滤和计算机模拟, 以解译大陆风化过程中 Li 同位素分馏的影响因素。然而, 目前所获得的结论观点不一, 乃至相互矛盾, 致使在运用 Li 同位素示踪大陆风化方面依然存在诸多争议和不确定性。

2.1 风化壳中 Li 同位素组成变化复杂

风化壳是岩石 / 矿物长期风化、雨水混合、次生矿物生成、地下水淋滤运移等过程共同作用下的产物 (Huh et al, 2002; Pistiner and Henderson, 2003; Ryu et al, 2014; Verney-Carron et al, 2015)。通过同一地区风化壳中的 Li 同位素分析发现, 不同风化壳的 Li 含量及其同位素表现出复杂的变化, 难以相互类比。例如, Huh et al (2002) 研究发现, 夏威夷地区玄武岩成因土壤中的 Li 含量是新鲜玄武岩的数倍, ⁷Li 值从剖面顶部到底部呈有规律的下降 (从 10‰ 降低到 3‰); 而 Pistiner and Henderson (2003) 的研究则发现该地区剖面土壤中⁷Li 值与新鲜玄武岩一致 (约 4‰); Ryu et al (2014) 的研究还发现, 从剖面浅部到深部, Li 元素含量逐步上升, 次生矿物 (尤其是高岭土) 的生成造成很大的 Li 同位素分馏。此外, 新鲜岩石初始风化时是否产生 Li 同位素分馏也没

有定论 (Pistiner and Henderson, 2003; Ryu et al, 2014)。尽管如此, 风化壳中的 Li 同位素研究取得了以下共识: (1) 风化壳形成过程中总伴随着外来物质的加入, 这在 Li 含量和⁷Li 值上都有反映 (Huh et al, 2002; Pistiner and Henderson, 2003; Kisakürek et al, 2004b; Liu et al, 2013; Ryu et al, 2014); (2)⁶Li 易被粘土吸附, 而⁷Li 易进入流体的性质可用来示踪古地下水水位 (Kisakürek et al, 2004b; Rudnick et al, 2004); (3) 次生矿物的生成和溶解是风化壳中⁷Li 值的重要制约要素之一 (Lemarchand et al, 2010; Liu et al, 2013; Ryu et al, 2014; Verney-Carron et al, 2015)。

风化壳形成于一个开放的体系, 其物质组成受母岩、气候、水文、生物作用、地形、形成时间、外来物质加入等要素共同制约 (Ollier, 1988; 李德文等, 2002)。因此, 不同风化壳之间化学组成变化的主导因素可能各不相同, 从而造成不同剖面得到不同的认识。

与其他风化壳相比, 具有均匀和稳定组成的黄土可能是开展 Li 同位素示踪大陆风化过程的理想对象。然而, 目前针对黄土 Li 同位素示踪风化过程的研究工作仅有一例。Tsai et al (2014) 对渭南黄土 - 古土壤序列中的 Li 同位素研究发现, 黄土中碎屑物质的 Li 含量和⁷Li 值变化不大, 与物源及矿物相关, 且与磁化率和粒径对应较好; 黄土中碳酸盐岩的 Li 含量和⁷Li 值变化则比较大, 并且其⁷Li 值与化学风化指数具有较好的相关性。

2.2 水系流域 Li 同位素的控制要素颇具争议

水系流域中的 Li 元素主要来源于岩石风化、人类活动和大气降水等 (Clergue et al, 2015; Dilling et al, 2015; Liu et al, 2015; Wang et al, 2015; Pogge von Strandmann et al, 2016, 2017)。目前, 各大陆主要流域河水和悬浮物 Li 同位素组成均已有数据报道 (图 2)。河水中 Li 的含量变化巨大, 为 0.06—81.2 μg·L⁻¹, 其⁷Li 值变化大约在 0.8‰—45.1‰, 全球河水平均⁷Li 值为 23‰ (Huh et al, 1998; Liu and Rudnick, 2011; Misra and Froelich, 2012)。河流悬浮物的⁷Li 组成大约在 -6.8‰—9.5‰, 总是低于各自河水中溶解态的⁷Li 值。然而, 由于大流域河水的 Li 同位素组成所受影响因素复杂, 主控因素各不相同, 很多问题还存在争议。近几

年,更多的研究转向对小流域Li收支的系统研究,得出了一些重要的认识(Lemarchand et al, 2010; Clergue et al, 2015),为Li同位素示踪大陆风化注入了活力。

河水中Li的来源及河水 $\delta^7\text{Li}$ 值的控制要素是目前运用Li同位素示踪大陆风化的争议焦点之一。以Huh等人为代表的观点认为,河水中的Li主要来源于硅酸盐岩的风化,其含量主要反映流域硅酸盐岩风化量;控制河水 $\delta^7\text{Li}$ 值的最主要因素不是岩性,而是风化过程中原生矿物的溶解与粘土矿物的吸附引起的同位素分馏之间的平衡,因此 $\delta^7\text{Li}$ 值反映流域硅酸盐岩的风化强度(Huh et al, 1998, 2001; Kisakürek et al, 2005; Vigier et al, 2009; Liu and Rudnick, 2011; Misra and Froelich, 2012; Froelich and Misra, 2014; Dellinger et al, 2014, 2015; Pogge von Strandmann et al, 2016)。与此同时,他们提出河水 $\delta^7\text{Li}$ 值变化还与以下过程有关:

(1)次生矿物的形成造成河水高的 $\delta^7\text{Li}$ 值(Vigier et al, 2009; Li and West, 2014);(2)冰川作用对河水的 $\delta^7\text{Li}$ 值影响微弱(Wimpenny et al, 2010b);(3)河水的pH对河水 $\delta^7\text{Li}$ 值有一定影响(Dellinger et al, 2015);(4)气候条件可能影响Li的元素行为,具体表现为寒冷气候环境下Li吸附在矿物表面且不发生分馏,而在温暖湿润气候条件下, ^{6}Li 优先进入次生矿物晶格且发生分馏(Godfrey et al, 2013)。然而,更多的研究表明,受水文循环、次生矿物溶解等因素影响,河水 $\delta^7\text{Li}$ 值与气候带关系不显著(Millot et al, 2010c; Liu et al, 2015; Pogge von Strandmann et al, 2016)。此外,Dellinger et al(2015)提出 $\delta^7\text{Li}$ 值变化与大陆风化之间非线性的关系。

与Huh等人的观点不同,以Pogge von Strandmann等人为代表的观点则认为,河水中Li的来源受多种因素控制, $\delta^7\text{Li}$ 值并不简单反映流域硅酸盐岩风化强度(Yoon, 2009; 罗超和郑洪波, 2011; Liu et al, 2011; Henchiri et al, 2014; Wang et al, 2015; Pogge von Strandmann and Henderson, 2015, Pogge von Strandmann et al, 2017)。例如:蒸发岩、火山活动带出的热液Li等均会较大程度地影响河流中的Li含量(Yoon, 2009; Henchiri et al, 2014; Wang et al, 2015);蒸发岩则可以影响河水的 $\delta^7\text{Li}$ 值(Yoon, 2009; Wang et al, 2015);悬浮物中不同粘土矿物和流

体之间不同的分馏系数可能是控制某些流域河水的 $\delta^7\text{Li}$ 值变化的主因(Yoon, 2009; Wang et al, 2015; Pogge von Strandmann et al, 2017)。由此,Pogge von Strandmann and Henderson(2015)提出了与Huh等截然相反的观点,认为河水 $\delta^7\text{Li}$ 值的变化反映的是冲积平原的形成及更多次生矿物的生成,而非化学风化强度。

那么,硅酸盐岩风化是不是控制河水Li同位素变化的主因呢?面对大河体系中Li含量及同位素组成的不确定性和多解性,近年来研究者转而开始研究小尺度水系流域(Catchment)中的Li收支。因为小尺度水系流域具有单一或稳定的岩性、相近的气候,这一研究思路展现出解决争议的巨大潜力,并已成为大陆风化研究的热点。通过小流域地球关键带(Earth's critical zone)的研究发现,在流域尺度上,Li同位素没有生命效应的重要结论(Lemarchand et al, 2010; Clergue et al, 2015; Pogge von Strandmann et al, 2016)。Li收支定量研究结果则显示,降雨、大气尘降对小流域Li收支有影响,其中对安山岩流域影响很大(Clergue et al, 2015),而对花岗岩质流域影响则很小(Lemarchand et al, 2010)。这可能主要是由于不同类型岩石Li含量之间数量级差别的“本底”效应造成的。目前,众多针对小流域的Li同位素研究工作正在进行中。

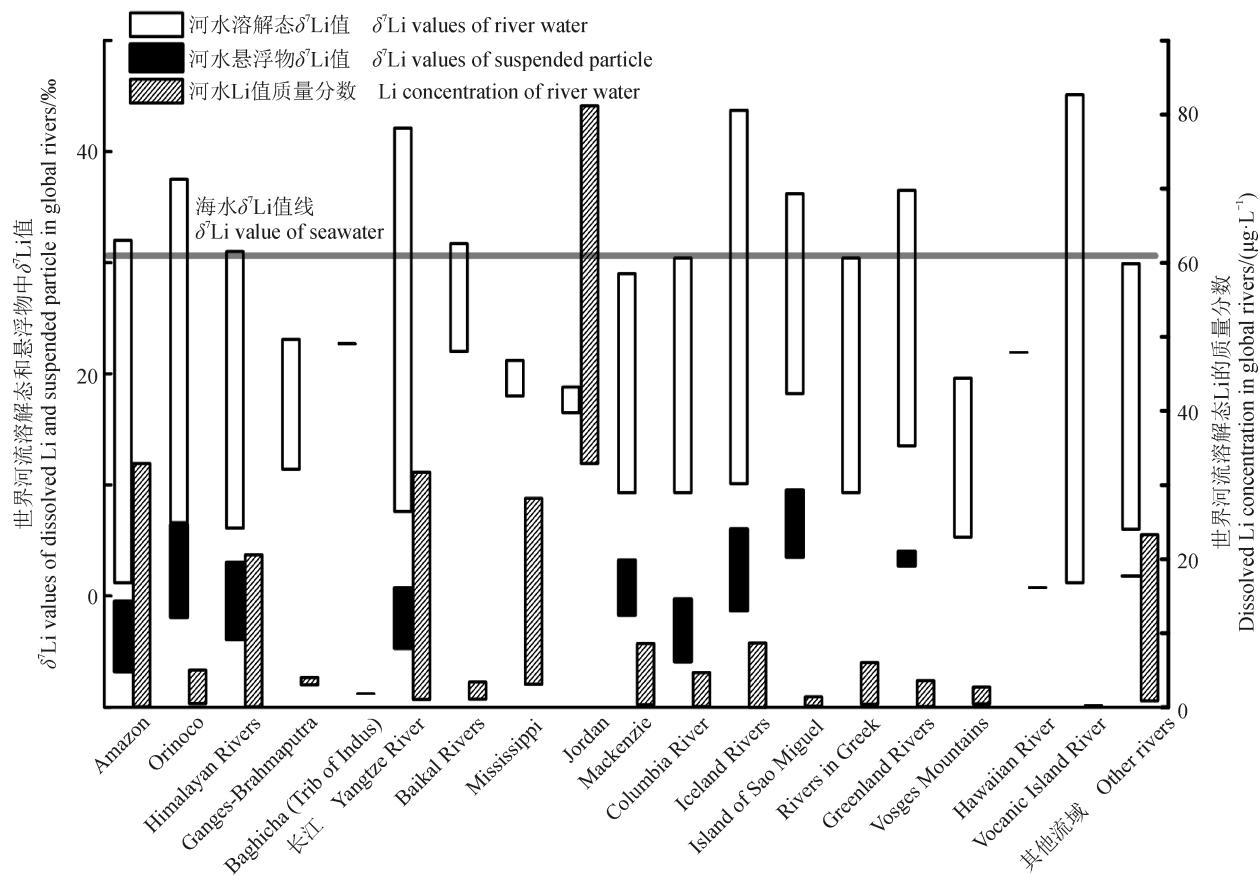
结合对大/小尺度水系流域中的Li同位素研究,大流域河水的Li含量及同位素组成的差异是各种制约要素的共同结果(Misra and Froelich, 2012; Wang et al, 2015)。为了确定全球尺度上不同制约要素之间的关系及其贡献大小,有必要开展更多小尺度流域水系中Li含量及其同位素组成的时空变化。开展小流域内高分辨率季节性河水化学与Li同位素结合的研究,或成为解决制约Li同位素变化控制要素争议的一个有效途径。

2.3 岩石淋滤实验尚无定论

Li同位素分馏的室内淋滤实验主要采用一些代表性的岩浆岩开展,但尚未获得一致的认识。例如:有的实验结果显示,基性岩石在酸性非平衡条件下淋滤时,没有发生Li同位素分馏(Wimpenny et al, 2010a),而在相同条件下酸性岩石淋滤时则发生Li同位素分馏(Pistiner and Henderson, 2003; Wunder et al, 2005; Millot et al, 2010b)。在近平衡条件下中性岩石的溶解实验中,次生矿

物形成时 Li 的类质同象被认为是造成 Li 同位素分馏的最主要因素 (Vigier et al, 2008; Wimpenny et al, 2010a; Millot et al, 2010b)。由于 Li 是中等不相容性元素, Li 在各类型岩石中不同的赋存状态可能是造成观察到上述现象的主要因素。在岩浆分异结晶过程中, Li 元素的不相容性使酸性岩具有高的 Li 含量。重要的是, 与基性岩中 Li 以类

质同象方式占据在稳定的晶格位置不同, 酸性岩中高 Li 含量可能使矿物晶体处于高能态, 酸性岩发生风化时 Li 优先被淋滤, 而 ${}^7\text{Li}$ 与水优先结合可使水分子处于低能态, 从而发生 Li 同位素分馏 (Huh et al, 2001); 相反, 基性岩的淋滤可能是矿物整体溶解的过程, 在没有次生矿物形成的情况下, 不发生同位素分馏。



数据来源: Amazon 引自 Dellinger et al (2015); Orinoco 引自 Huh et al (1998, 2001); Himalayan Rivers 引自 Kisakürek et al (2005); Ganges-Brahmaputra 引自 Huh et al (1998, 2001); Baghicha (Tributary of Indus) 引自 Huh et al (2001); 长江引自 Wang et al (2015); Baikal Rivers 引自 Huh et al (1998); Mississippi 引自 Chan et al (1992) 和 Huh et al (1998); Jordan 引自 Chan et al (1992); Mackenzie 引自 Huh et al (1998) 和 Millot et al (2010b); Columbia River 引自 Huh et al (1998) 和 Liu et al (2013); Iceland Rivers 引自 Pogge von Strandmann et al (2006) 和 Vigier et al (2009); Island of Sao Miguel 引自 Pogge von Strandmann et al (2004a); Rivers in Greek 引自 Liu and Rudnick (2011); Greenland Rivers 引自 Wimpenny et al (2010b); Vosges Mountains 引自 Lemarchand et al (2010); Hawaiian River 引自 Pistiner and Henderson (2003); Volcanic Island River 引自 Henchiri et al (2014); 其他流域数据引自 Huh et al (1998, 2001)。
Amazon data are from Dellinger et al (2015); Orinoco data from Huh et al (1998, 2001); Himalayan Rivers data from Kisakürek et al (2005); Ganges-Brahmaputra data from Huh et al (1998, 2001); Baghicha (Tributary of Indus) data from Huh et al (2001); Yangtze River data from Wang et al (2015); Baikal Rivers data from Huh et al (1998); Mississippi data from Chan et al (1992) and Huh et al (1998); Jordan data from Chan et al (1992); Mackenzie data from Huh et al (1998) and Millot et al (2010b); Columbia River data from Huh et al (1998) and Liu et al (2013); Iceland Rivers data from Pogge von Strandmann et al (2006) and Vigier et al (2009); Island of Sao Miguel data from Pogge von Strandmann et al (2004a); Rivers in Greek data from Liu and Rudnick (2011); Greenland Rivers data from Wimpenny et al (2010b); Vosges Mountains catchment data from Lemarchand et al (2010); Hawaiian River data from Pistiner and Henderson (2003); Volcanic Island River data from Henchiri et al (2014); and other rivers data from Huh et al (1998, 2001).

图 2 已报道的全球主要河流溶解态和悬浮物中 $\delta^7\text{Li}$ 值和 Li 质量分数的变化范围

Fig.2 Concentrations and isotopic compositions of lithium reported in dissolved and suspended phases in global major rivers

此外, Li 吸附到矿物表面是否引起 Li 同位素分馏也还没有定论, 有实验发现吸附会影响 Li 的分馏, 且分馏程度与所吸附的矿物类型有关 (Pistiner and Henderson, 2003), 有些则认为该过程不发生分馏 (罗超和郑洪波, 2011; Wimpenny et al, 2015)。同时, 相对于巨大的动力学过程引起的 Li 同位素分馏, 低温条件下 (地表环境) 扩散所致的平衡分馏效应基本上可忽略 (Verney-Carron et al, 2011)。

2.4 模型模拟尚处摸索阶段

现代计算机的强大计算功能为定量解决地学问题提供了新的渠道。上述岩石矿物 Li 同位素淋滤实验获得的基础数据为计算机模拟 Li 同位素体系创造了条件。目前有关 Li 同位素分馏的模拟结果并不多, 主要提出了多种控制河水 Li 同位素组成变化要素的组合, 例如气候对河水中的 Li 通量的作用 (Vigier and Goddériss, 2015)、次生矿物的溶解 (Bouchez et al, 2013)、粘土和土壤对河水 Li 的固定 (Bouchez et al, 2013; Li and West, 2014; Wanner et al, 2014) 等因素分别被不同研究者认为对水体 $\delta^7\text{Li}$ 值影响很大。由于上述制约 Li 同位素分馏的不平衡、非线性复杂过程, 模型模拟的计算结果由初始状态和边界条件决定, 因此实验淋滤结果的好坏直接决定模拟结论的正确与否, 反之模拟结论是否合理又需要水系流域、风化壳的研究来佐证。应该看到, 模型模拟拓展了 Li 同位素示踪大陆风化的研究手段, 但是室内淋滤数据依然十分有限, 这可能造成模拟结果一定的“蝴蝶效应”。

3 问题与挑战

上述可知, 对 Li 同位素示踪大陆风化的研究已取得一些重要成果, 在风化壳、水系流域、实验室模拟等方面均取得了重要进展, 这些进展极大地促进了对 Li 同位素在表生地球化学行为的认识以及其示踪大陆风化的可行性。事实上, 对 Li 同位素体系示踪大陆风化的研究始于检验“构造抬升 - 化学风化 - 气候变化”假说。该假说认为构造隆升造成大陆岩石风化加强进而导致了晚新生代气候变冷。此前 Sr、Nd、Os 等多个同位素体系参与验证该假说, 但因各自的局限性最终未能获得理想的结果。与这几个同位素体系类似, Li 同位素体系也面临着如何定量评估各影响因素, 进

而示踪大陆硅酸盐岩风化的困境。然而, Li 同位素巨大的分馏远比原始风化物质组成范围大得多, 只有风化过程才会引起这么大的分馏 (Lemarchand et al, 2010; Misra and Froelich, 2012; Sauzeat et al, 2015)。因此, 若不细分风化过程中各具体影响因素, 那么 Li 同位素分馏在一定程度上应该可以反映化学风化的变化。基于以上事实, Li 同位素的确加深了人们对大陆风化的理解, 但是要使 Li 同位素体系更有效地示踪大陆风化, 需要面对以下三大挑战:

(1) 如何扣除深部热液的影响。深部热液具有高 Li 含量, 由火山、地震等构造运动带入地表体系的深部热液可能成为 Li 同位素体系示踪硅酸盐岩风化的一个干扰项。因此, 在一些构造活跃的高山流域, 在运用 Li 同位素示踪硅酸盐岩风化强度时, 需考虑深部热液的 Li 贡献及其同位素组成。

(2) 如何量化风化过程中的 Li 同位素分馏的各个控制要素。目前已经认识到的制约 Li 同位素分馏的要素包括侵蚀、溶解、吸附、次生矿物的形成、pH 等 (Huh et al, 1998, 2001; Kisakürek et al, 2005; Vigier et al, 2009; Liu and Rudnick, 2011; Dellinger et al, 2014, 2015; Froelich and Misra, 2014), 这些因素对 $\delta^7\text{Li}$ 值的贡献大小还不能有效量化。由于形成演化过程和形成时间不同, 风化壳 Li 同位素特征难以进行类比, 因此风化壳剖面的研究恐难以量化 Li 同位素分馏的控制要素。相反, 对单一岩性流域的连续监测, 有望缩小制约要素, 进而减少 Li 同位素示踪风化的不确定性。

(3) 如何量化人为活动对 Li 同位素分馏的贡献。 ${}^6\text{Li}$ 作为重要的化工、电源和核聚变原料 (贾小波等, 2007; Arikawa et al, 2010), 目前已有多种人工方法富集分离 ${}^6\text{Li}$ (Tatenuma et al, 2001), 使 $\delta^7\text{Li}$ 值超过自然分馏的百倍以上 (Qi et al, 1997)。因此, 尽管对有些流域尚未构成影响 (Négrel et al, 2010; Wang et al, 2015), 但如何量化人类活动对水体 $\delta^7\text{Li}$ 值的贡献成为 Li 同位素示踪硅酸盐岩风化需面对的又一大难题 (Qi et al, 1997; Négrel et al, 2010)。

4 研究展望

地表环境下 Li 的地球化学行为使其同位素在

示踪大陆风化方面取得了一些新的认识和进展, 同时也面临着诸多挑战。Li 同位素体系能否作为大陆风化的示踪指标, 还需要加强以下四个方面 的研究:

(1) 细化岩石 / 矿物淋滤实验: 通过实验室控制实验条件, 从单一变量着手, 系统评价不同要素贡献, 为 Li 同位素体系示踪大陆风化提供可靠依据;

(2) 大小流域相结合: 对比研究单一岩性、相同气候条件下、没有人为输入的小流域内 Li 收支及其同位素行为, 结合大流域 Li 同位素的宏观特征, 将进一步明确 Li 同位素来源和分馏的制约要素。鉴于黄土具有相对均匀的、可代表上地壳的地球化学成分, 查明黄土现代风化过程中 Li 同位素行为有望对认识 Li 同位素示踪整体大陆风化产生不可替代的推动作用, 但这项工作目前尚未引起足够重视;

(3) 多同位素体系相互补充与验证: 地表风化作用是复杂的, Li 同位素获得的认识可能只是风化过程的某一侧面, 存在不确定性或多解性, 因此需要与其他同位素相互补充与验证, 以更全面地认识化学风化过程及其制约要素;

(4) 加强计算机模型模拟研究: 目前计算机模型模拟 Li 同位素体系的研究尚处于起步阶段, 所得认识还较为有限, 但计算机模拟将最终为 Li 同位素示踪大陆风化的定量研究提供最有益的帮助。

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