

土壤稀土元素的迁移-富集机制及其生态效应

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摘要: 稀土被称为“工业维生素”。随着稀土元素使用日益增加, 导致其在土壤广泛分布并不断积累。土壤中稀土元素的地球化学过程已成为全球关注的热点。本文通过检索土壤环境稀土元素方向研究文献, 综述土壤稀土元素迁移-富集机制的研究进展, 识别土壤稀土元素主要来源, 探讨土壤稀土元素的含量分布、分馏特征和赋存相态, 以及稀土元素迁移富集和分馏的影响因素, 分析稀土元素对土壤理化性质、动物、植物和微生物等产生的生态效应, 以及对人体健康的潜在威胁, 提出土壤稀土元素未来研究的关键问题和方向。相关认识有助于理解和掌握稀土元素在土壤中的迁移和归趋, 并为稀土污染防治提供理论依据。

关键词: 土壤; 稀土元素; 迁移-富集; 生态效应; 健康风险

Migration enrichment mechanism and ecological effects of rare earth elements in soil

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Abstract: *Background, aim, and scope* Rare earth elements (REEs) are known as “industrial vitamins”. The increasing use of REEs has led to their widespread distribution and accumulation in the soil, and they are considered emerging pollutants. Southern China contains granite bodies in which ion-adsorption-type REE deposits have formed due to surface weathering. REEs are highly soluble and prone to migrate by leaching into the surrounding soil environment, a process affected by the weathering crust medium, climatic conditions and geomorphic characteristics of the natural environment. REE pollution is believed to pose a long-term ecological risk in China. As both an important “source” and “sink” of pollutants, the soil environment is an important window through which to study the migration and enrichment processes of REEs. To make efficient use of REE resources, while controlling associated pollution, geochemists and environmental scientists worldwide have investigated the geochemical processes of REEs in soil. This paper presents an analysis of the sources, processes

收稿日期: 2022-03-24; 录用日期: 2022-06-28; 网络出版: 2022-07-13

Received Date: 2022-03-24; **Accepted Date:** 2022-06-28; **Online first:** 2022-07-13

基金项目: 国家自然科学基金项目(41771024)

Foundation Item: National Natural Science Foundation of China (41771024)

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引用格式: 林卓玲, 黄光庆. 2023. 土壤稀土元素的迁移-富集机制及其生态效应 [J]. 地球环境学报, 14(5): 521–538.

Citation: Lin Z L, Huang G Q. 2023. Migration enrichment mechanism and ecological effects of rare earth elements in soil [J]. *Journal of Earth Environment*, 14(5): 521–538.

and ecological risks relating to soil REE pollution based on a review of research on REEs in soil surface environments, and proposes future research and development directions and key problems relating to soil REEs. The paper will provide a reference to support further exploration of the migration and enrichment of REEs in soil and the evaluation of the associated ecological risk. **Materials and methods** Based on the literature on REEs in soil environments in China and beyond, this paper summarizes the progress of research into the migration, enrichment mechanism and ecological effects of REEs in soil. **Results** The main sources of REEs in soil are the natural leaching of ion-adsorption-type REE ores, mining of REEs, agricultural application of fertilizer, and vehicular and dust emissions. Differences in soil texture and utilization type lead to differences in the content and spatial distribution of REEs. Generally, REEs exist in either an inorganic-bound state, an organic-bound state or an iron manganese oxide-bound state. Depending on the pH value, redox conditions, soil texture and other factors, REEs undergo migration and enrichment at a range of scales. These processes have an adverse impact on the physical and chemical properties of soil and on animals, plants and microorganisms, and pose varying degrees of risk to human health through respiratory exposure and the food chain. Studies have revealed the potential health risks of REEs at the molecular, cellular and individual levels. **Discussion** The migration and enrichment of REEs are restricted by multiple factors. The occurrence phases of REEs in soil affect their content, migration behavior and bioavailability. Specifically, REEs form complexes with inorganic oxyacid anions and organic ligands, and are mainly distributed in the soil in the form of inorganic-bound, organic-bound or iron manganese oxide-bound phases. The distribution of these phases results in marked differences in the content and migration rate of REEs in the soil. Meanwhile, the migration, enrichment and fractionation of REEs in soil can be understood as a series of complex chemical processes, including weathering-leaching, oxidation-reduction, dissolution-precipitation, adsorption-desorption and others. These processes are affected by the soil pH value, redox conditions, soil texture and other factors, resulting in variations in the migration and enrichment of REEs and corresponding changes in their geochemically driven fractionation behavior. REEs differ in their abundance distribution patterns and environmental migration potential. Future research is expected to focus on further tracking and investigating the sources and contributors of REEs in soil from the perspective of multiple temporal and spatial scales and land-use types, revealing the dynamic mechanism of their release and migration, evaluating the dose-biological response behavior of REE compounds as pollutants and the corresponding health risks, exploring the response strategies of microorganisms to REE-rich environments and developing new technologies for resource utilization of REEs. **Conclusions** Through field monitoring, experimental simulations and theoretical model calculations, researchers have conducted in-depth investigations of the abundance, differential activation and migration mechanisms, morphology and influencing factors of REEs. Due to the complexity of soil pollution profiles and the diversity of sources of REEs, the behavior of REEs in soil and the associated risks are not yet clearly understood. The differentiation and occurrence forms of REEs significantly affect their migration process and environmental effects in the soil, in turn affecting the biological, ecological and human health risks these elements pose. **Recommendations and perspectives** Relevant knowledge is helpful for understanding and managing the migration and fate of REEs in soil, and providing a theoretical basis for the prevention and control of REE pollution.

Key words: soil; rare earth elements; migration and enrichment; ecological effect; health risk

稀土元素（rare earth elements, REEs）是化学元素周期表中镧系 15 种元素和钪（Sc）、钇（Y）元素的总称。稀土元素被称为“工业维生素”，在汽车制造、电子工业、可再生能源、国防军工等领域具有重要应用价值（Gwenzi et al.,

2018; Balaram, 2019），已成为全球公认的重要战略资源。我国已探明的稀土资源储存量位居世界首位，约占世界总储量的 33%，年产量占世界总产量的 85% 以上（王猛等, 2019）。离子吸附型稀土矿床是花岗岩和花岗斑岩（稀土元素相对

丰度较高)在化学风化作用下分解溶解, 随后稀土矿物溶液迁移渗透过程中被黏土矿物吸附和富集而形成的(Yang et al., 2013)。受自然环境中风化壳介质、气候条件和地貌特征影响, 稀土元素极易受淋滤作用发生溶出和迁移, 进入周边土壤中。在稀土开采和选冶过程中, 稀土元素、重金属和浸矿剂也对矿区周边土壤产生较为严重的污染, 导致土壤环境质量下降, 严重威胁周边居民健康(Dutta et al., 2016)。另外, 长期施用含稀土的肥料以及对电子废弃物的不当处理, 都会导致稀土元素在土壤环境中不断积累(Ramos et al., 2016)。当前, 稀土元素被认为是新兴污染物, 预计在未来相当长时间内将对相关地区的生态环境造成危害。

作为重要的“源”和“汇”, 土壤是稀土

元素迁移-富集的重要归趋。国内外研究通过野外监测、实验模拟和理论模型计算, 对土壤中稀土元素的浓度水平、活化与迁移机制、赋存相态及其影响因素等多个方面进行了深入系统研究。本文总结有关土壤中稀土污染的相关研究进展, 分析土壤稀土元素来源、迁移-富集影响因素及生态风险, 提出土壤稀土污染防治未来研究的方向与重点, 以期为掌握土壤稀土元素迁移富集规律与生态环境风险评估提供科学依据。

1 土壤稀土元素的来源及途径

土壤中稀土元素主要来源于离子吸附型稀土矿的自然淋滤作用、稀土开采、农业肥料施用、交通道路及灰尘等排放途径(图1)。

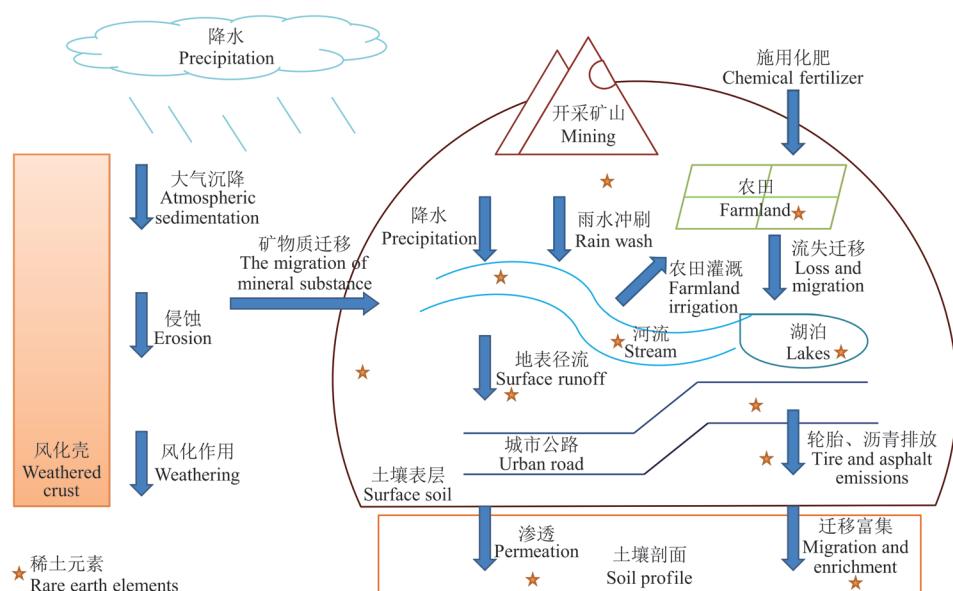


图1 土壤中稀土元素的来源与途径
Fig. 1 Sources and pathways of rare earth elements entering into soil

1.1 离子吸附型稀土矿床的自然淋滤作用

离子吸附型稀土矿床广泛分布于全球亚热带-热带地区, 包括我国南方七省, 以及越南、老挝、缅甸、泰国、菲律宾和马达加斯加等国家/地区。离子吸附型稀土矿床中的稀土离子主要富集在全风化层, 稀土元素主要以三价阳离子形态存在(占稀土元素总量的60%—95%), 吸附在高岭石和埃洛石等次生矿物表面(Liu et al., 2019; Borst et al., 2020; Li and Zhou, 2020)。风化壳上层部位受到化学风化作用最强烈, 从上层部位到底

部化学风化强度逐渐减弱(谢明君等, 2022)。在降雨、地表径流以及地下水作用下, 加之土壤微生物代谢活动释放有机酸, 以及有机质分解为腐殖酸, 风化壳呈弱酸性(pH值约5—6.5)(Li et al., 2019)。该酸性条件非常利于原岩及风化壳中的稀土元素迁移进入土壤、河流和地下水中, 造成周边环境稀土含量异常。

1.2 稀土开采排放

稀土开采活动更进一步促使稀土离子活化失稳, 向周围环境迁移扩散, 导致稀土矿区周边土

壤的稀土浓度比母岩高得多，特别是在离子吸附型稀土矿床开采过程中稀土元素向周边土壤迁移富集的现象更为显著。目前稀土矿床开采主要采用原位浸取法，即利用浸取剂硫酸铵溶液中铵离子的离子交换作用，提取浸出离子吸附态稀土元素，同时硫酸根与稀土元素产生络合反应，加速稀土元素的释放和迁移。由于风化壳富含黏土矿物易于吸附稀土元素，且矿床疏松、孔隙网复杂，浸取电解质和浸出液随着径流和壤中流等扩散至周边环境，并继续与下游沉积物、风化壳和土壤中的稀土离子进行交换，加速稀土活化和迁移。而大部分稀土残留在土壤表层，随后通过垂直迁移逐渐富集到土壤深层部位以及和地下水系统，在雨水淋溶、冲刷和地表径流等作用下迁移至周边湖泊、河流等地表水系统。因此，离子吸附型稀土矿开采对矿区及周边地区土壤产生了重要影响，稀土元素含量范围在 $396\text{--}2314\text{ mg}\cdot\text{kg}^{-1}$ ，其最低浓度也是我国土壤背景值的两倍（金姝兰等，2014；Liang et al., 2014）。稀土开采活动不仅导致废弃尾矿的土壤稀土元素含量高，在尾矿土壤、周边环境的水体、沉积物和农田土壤中也发现了稀土富集（Liu et al., 2019），土壤稀土含量与矿区距离呈显著负相关关系（王学锋等，2019）。

1.3 农业肥料排放

稀土元素可作为微肥料广泛应用于农业，提高农产品质量和产量（Gueroult et al., 2018）。比如：磷灰石是重要的富稀土矿物，可用于制造洗涤剂、动物饲料和磷酸盐（ PO_4^{3-} ）肥料，其长期使用导致土壤稀土元素浓度升高。据报道，农业土壤稀土元素含量范围为 $16\text{--}2450\text{ mg}\cdot\text{kg}^{-1}$ （Mihajlovic and Rinklebe, 2018；Adeel et al., 2019；Mihajlovic et al., 2019），其中 La_2O_3 平均含量为 $6.6\text{--}50\text{ mg}\cdot\text{kg}^{-1}$ （Li et al., 2018）。泰国地区由于水稻种植长期施用磷肥，引起土壤中稀土的积累（Sukitprapanon et al., 2019）。在我国、德国、伊朗、马其顿和葡萄牙等国家的土壤中也存在类似现象（Zhang et al., 2006；Mihajlovic et al., 2014；Mazhari and Sharifiyan Attar, 2015；Brito et al., 2018）。

1.4 交通道路及灰尘排放

大气中稀土元素沉降于表土上，也是土壤稀

土元素来源之一（Mihajlovic et al., 2019）。废气中稀土元素来源于汽油发动机催化转化器中使用的稀土元素（Borowiak et al., 2018；Yuan et al., 2018）。研究发现，公路边土壤比周边地区更富含稀土元素（Borowiak et al., 2018），尤其是 Nd 和 Ce。但随着道路距离增加，大多数稀土元素普遍减少（Mleczeck et al., 2018）。Mleczeck et al. (2021) 指出公路边土壤稀土元素含量为 $24\text{--}78\text{ mg}\cdot\text{kg}^{-1}$ ，与交通强度密切相关。轮胎、沥青释放稀土元素到环境中，也是公路边土壤稀土污染的主要来源。在稀土工业城市，道路灰尘含有大量的轻稀土元素（Li et al., 2020），扬尘粒径越细，其稀土元素含量和污染程度越高，促进了周边土壤稀土元素积累。

2 土壤稀土元素含量与分馏特征

与稀土矿区土壤相比，关于非矿区土壤中稀土元素浓度水平的报道较少。不同土壤质地、利用类型均会造成土壤稀土元素含量和分馏特征存在差异性。目前，国内外学者从不同空间尺度研究了土壤中稀土元素的含量水平分布和分馏特征。

2.1 土壤稀土元素含量水平分布

关于稀土矿区中稀土元素的污染，如表 1 所示，Liu et al. (2019) 关注了世界上最大离子吸附型稀土矿（江西赣州）周边环境土壤，发现稀土在尾矿土和农田土壤中平均含量分别为 $392\text{ mg}\cdot\text{kg}^{-1}$ 和 $928\text{ mg}\cdot\text{kg}^{-1}$ ，揭示了稀土元素在土壤-水体-底泥等不同环境介质中的分布差异和联系，指出离子吸附型稀土矿可能对周围环境构成威胁。关于非矿区土壤，Huang et al. (2019) 研究了福建九龙河流域水稻土壤，发现稀土的平均总浓度为 $267\text{ mg}\cdot\text{kg}^{-1}$ ，空间上各元素浓度呈均匀分布。在珠江三角洲地区，冲积层土壤的稀土含量在 $37\text{--}550\text{ mg}\cdot\text{kg}^{-1}$ ，平均值为 $268\text{ mg}\cdot\text{kg}^{-1}$ (Chang et al., 2016)。以上稀土含量均值都高于我国土壤稀土元素背景值 ($187.6\text{ mg}\cdot\text{kg}^{-1}$) (魏复盛等, 1991)。

其他国家稀土富集区土壤中稀土元素的含量水平和分布也相继受到关注。泰国中部低平原长期栽培水稻，表层土壤 (0—20 cm) 中稀土元素含量最高，稀土元素含量从酸性硫酸盐土壤层逐渐向表层增加（Sukitprapanon et al., 2019）。在巴西亚马逊河的两个水文流域（索里莫斯和里约

热内格罗) 以及城市土壤中, 索里莫斯、里约热内格罗盆地和城市区域土壤表层及次表层的稀土元素含量比较接近, 索里莫斯土壤稀土元素含量明显高于里约热内格罗盆地和城市区域, 索里莫

斯土壤中的稀土浓度范围参考值为 $0.01 \text{ mg} \cdot \text{kg}^{-1}$ (Lu) 至 $145.6 \text{ mg} \cdot \text{kg}^{-1}$ (Ce), 里约热内卢土壤中为 $0.05 \text{ mg} \cdot \text{kg}^{-1}$ (Lu) 至 $15.8 \text{ mg} \cdot \text{kg}^{-1}$ (Ce) (da Silva Ferreira et al., 2021)。

表 1 土壤稀土元素含量分布 (离子吸附型稀土富集区)
Tab. 1 Content distribution of rare earth elements in soil (ion-adsorption rare earth enrichment area)

国家 / 地区 Country / Region	含量范围 Content range	分馏情况 Fractionation	参考文献 References
我国江西省 赣州市 Ganzhou City, Jiangxi Province, China	废弃 6 a、9 a 和 14 a 的尾矿土壤中 REEs 的含量分别为 $187\text{--}1148 \text{ mg} \cdot \text{kg}^{-1}$ 、 $240\text{--}456 \text{ mg} \cdot \text{kg}^{-1}$ 和 $315\text{--}485 \text{ mg} \cdot \text{kg}^{-1}$, 均值为 $(392\pm225) \text{ mg} \cdot \text{kg}^{-1}$ 。受污染农田采样点土壤中 REEs 含量分别为 $(760\pm85.4) \text{ mg} \cdot \text{kg}^{-1}$ 、 $(1100\pm26.9) \text{ mg} \cdot \text{kg}^{-1}$ 。 The contents of REEs in tailings soil abandoned for 6, 9 and 14 years were $187\text{--}1148 \text{ mg} \cdot \text{kg}^{-1}$, $240\text{--}456 \text{ mg} \cdot \text{kg}^{-1}$ and $315\text{--}485 \text{ mg} \cdot \text{kg}^{-1}$, respectively, with an average of $(392\pm225) \text{ mg} \cdot \text{kg}^{-1}$. The contents of REEs in the soil of contaminated farmland sampling points were $(760\pm85.4) \text{ mg} \cdot \text{kg}^{-1}$ and $(1100\pm26.9) \text{ mg} \cdot \text{kg}^{-1}$, respectively.	多数尾矿土壤均表现出中、重稀土元素富集。多数尾矿土表现为正 Ce 异常。尾矿土无明显的 Eu 异常。农田土壤 HREEs 元素富集。农田土壤中稀土元素污染主要是由废弃尾矿淋溶作用引起。 Most tailings soils showed enrichment of MREEs and HREEs. Most tailings soils show positive Ce anomaly. There is no obvious Eu anomaly in tailings soil. The pollution of REEs in farmland soil was mainly caused by the leaching of waste tailings.	Liu et al., 2019
我国广东省 珠三角地区 Pearl River Delta, Guangdong Province, China	表层土壤样品中 $\sum \text{REEs}$ 、 $\sum \text{LREEs}$ 和 $\sum \text{HREEs}$ 分别为 $37\text{--}550 \text{ mg} \cdot \text{kg}^{-1}$ 、 $32\text{--}480 \text{ mg} \cdot \text{kg}^{-1}$ 和 $4.8\text{--}128 \text{ mg} \cdot \text{kg}^{-1}$ 。 $\sum \text{Ce}$ 约占 80%, $\sum \text{Y}$ 约占 20%。 $\sum \text{REEs}$ 、 $\sum \text{LREEs}$ 和 $\sum \text{HREEs}$ were $37\text{--}550 \text{ mg} \cdot \text{kg}^{-1}$, $32\text{--}480 \text{ mg} \cdot \text{kg}^{-1}$ and $4.8\text{--}128 \text{ mg} \cdot \text{kg}^{-1}$, respectively. $\sum \text{Ce}$ accounts for about 80%, $\sum \text{Y}$ accounts for about 20%.	多数土壤具有轻微 Ce 正异常和强烈 Eu 负异常。 Most soils exhibited slight positive Ce anomaly and strong negative Eu anomaly.	Chang et al., 2016
我国贵州省 喀斯特地区 Karst areas in Guizhou Province, China	风化剖面中 $\sum \text{REEs}$ 为 $167.4\text{--}1814.2 \text{ mg} \cdot \text{kg}^{-1}$, $\sum \text{LREEs}$ 和 $\sum \text{HREEs}$ 分别为 $137.1\text{--}1129.7 \text{ mg} \cdot \text{kg}^{-1}$ 、 $30.29\text{--}684.49 \text{ mg} \cdot \text{kg}^{-1}$ 。 $\sum \text{REEs}$ in weathering profile was $167.4\text{--}1814.2 \text{ mg} \cdot \text{kg}^{-1}$, $\sum \text{LREEs}$ and $\sum \text{HREEs}$ were $137.1\text{--}1129.7 \text{ mg} \cdot \text{kg}^{-1}$ and $30.29\text{--}684.49 \text{ mg} \cdot \text{kg}^{-1}$, respectively.	REE 含量从剖面底部到表层土呈下降趋势, 稀土元素分馏受到持续风化作用的影响。高 LREEs/HREEs 和 $(\text{La}/\text{Yb})_N$ 比值证实土壤成土过程中 LREEs 比 HREEs 更具流动性。 REEs content decreased from the bottom of the profile to the topsoil. REE fractionation was affected by continuous weathering. High LREEs/HREEs and $(\text{La}/\text{Yb})_N$ ratios confirmed that LREEs was more mobile than HREEs in the process of soil formation.	Chang et al., 2019
我国福建省 九龙流域 Jiulong River Basin, Fujian Province, China	稀土元素平均总浓度为 $266.86 \text{ mg} \cdot \text{kg}^{-1}$, Ce 均值达 $112.1 \text{ mg} \cdot \text{kg}^{-1}$ (占稀土总量的 42.3%), $\sum \text{LREEs}$ 、 $\sum \text{HREEs}$ 的平均值分别为 $239.45 \text{ mg} \cdot \text{kg}^{-1}$ 、 $27.42 \text{ mg} \cdot \text{kg}^{-1}$ 。 The average total concentration of REEs was $266.86 \text{ mg} \cdot \text{kg}^{-1}$, the mean value of Ce was $112.1 \text{ mg} \cdot \text{kg}^{-1}$ (accounting for 42.3% of the total rare earth), and the mean values of $\sum \text{LREEs}$ and $\sum \text{HREEs}$ were $239.45 \text{ mg} \cdot \text{kg}^{-1}$ and $27.42 \text{ mg} \cdot \text{kg}^{-1}$.	水稻土中稀土元素的分布格局为 LREEs 富集型, LREEs 表现出明显的分馏, HREEs 没有明显的分馏。土壤中 Eu 呈显著负异常, Ce 无显著异常。土壤中稀土元素均略有富集。 The distribution pattern of REEs in paddy soil was LREEs enrichment type, LREEs showed obvious differentiation, and HREEs had no obvious differentiation. Eu in soil showed significant negative anomaly, and Ce had no significant anomaly. REEs in soil were slightly enriched.	Huang et al., 2019
泰国中部低平原 热带地区 Central Thailand low plain tropical region	Sc、Y、La、Ce、Nd 和 Gd 含量分别为 $5.8\text{--}12 \text{ mg} \cdot \text{kg}^{-1}$ 、 $5.0\text{--}34 \text{ mg} \cdot \text{kg}^{-1}$ 、 $13\text{--}36 \text{ mg} \cdot \text{kg}^{-1}$ 、 $28\text{--}81 \text{ mg} \cdot \text{kg}^{-1}$ 、 $11\text{--}36 \text{ mg} \cdot \text{kg}^{-1}$ 和 $1.0\text{--}7.3 \text{ mg} \cdot \text{kg}^{-1}$ 。 The contents of Sc, Y, La, Ce, Nd and Gd were $5.8\text{--}12 \text{ mg} \cdot \text{kg}^{-1}$, $5.0\text{--}34 \text{ mg} \cdot \text{kg}^{-1}$, $13\text{--}36 \text{ mg} \cdot \text{kg}^{-1}$, $28\text{--}81 \text{ mg} \cdot \text{kg}^{-1}$, $11\text{--}36 \text{ mg} \cdot \text{kg}^{-1}$ and $1.0\text{--}7.3 \text{ mg} \cdot \text{kg}^{-1}$, respectively.	总稀土元素的最小浓度出现在硫酸盐层的底部。土壤中 REE 的最高浓度出现在表土中。 The minimum concentration of REEs appeared at the bottom of the sulfate layer. The highest concentration of REEs in soil was observed in topsoil.	Sukitprapanon et al., 2019

(待续 To be continued)

(续表 1 Continued Tab. 1)

国家 / 地区 Country / Region	含量范围 Content range	分馏情况 Fractionation	参考文献 References
巴西亚马逊河 流域地区 Brazilian Amazon region	索里莫斯土壤表层和次表层的 REEs 均值为 $109.28 \text{ mg} \cdot \text{kg}^{-1}$ 和 $94.11 \text{ mg} \cdot \text{kg}^{-1}$ ，LREEs 含量均值为 $78.09 \text{ mg} \cdot \text{kg}^{-1}$ 和 $72.01 \text{ mg} \cdot \text{kg}^{-1}$ ，HREEs 含量均值为 $4.85 \text{ mg} \cdot \text{kg}^{-1}$ 和 $4.60 \text{ mg} \cdot \text{kg}^{-1}$ 。里约热内格罗盆地土壤表层和次表层稀土元素含量分别为 $43.9515 \text{ mg} \cdot \text{kg}^{-1}$ 和 $38.40 \text{ mg} \cdot \text{kg}^{-1}$ ，LREEs 含量均值为 $33.26 \text{ mg} \cdot \text{kg}^{-1}$ ($0-20 \text{ cm}$) 和 $29.18 \text{ mg} \cdot \text{kg}^{-1}$ ，HREEs 含量均值为 $2.07 \text{ mg} \cdot \text{kg}^{-1}$ 和 $1.68 \text{ mg} \cdot \text{kg}^{-1}$ 。城市表土 REEs、LREEs 和 HREEs 分别为 $38.62 \text{ mg} \cdot \text{kg}^{-1}$ 、 $32.44 \text{ mg} \cdot \text{kg}^{-1}$ 、 $1.47 \text{ mg} \cdot \text{kg}^{-1}$ 。	受到地质特征、沉积物和人为输入之间动态相互作用影响，土壤中稀土元素含量呈现出很大的变化，轻重稀土元素分馏明显，LREEs 相对于 HREEs 富集。	da Silva Ferreira et al., 2021

2.2 土壤环境中稀土元素分馏特征

各稀土元素在地壳中的平均丰度相差很大 (Borst et al., 2020; Huang et al., 2021)，一般符合 Oddo-Harkins 规则，稀土元素含量随原子序数增加而减少，原子序数为偶数的稀土元素含量高于相邻原子序数为奇数的含量。通过将所测样品的稀土元素含量值标准化后，按原子序数绘制稀土元素的变化曲线，来阐释稀土元素配分模式 (Laveuf and Cornu, 2009)。稀土元素测量值标准化常用的参照数据包括：球粒陨石 (chondrite)、北美页岩 (NASC)、后太古代澳大利亚页岩 (PAAS)、上层地壳 (UCC) (Godwyn-Paulson et al., 2022)。

由于稀土元素电子结构具有 4f 亚层电子数目逐渐递增的规律，化学性质发生规律性改变，如：(1) 与轻稀土相比，重稀土更容易水解；(2) 相较于轻稀土，重稀土优先与环境中的无机配位体 CO_3^{2-} 、 SO_4^{2-} 等形成络合物，在环境中迁移能力更强；(3) 镥 (Ce) 和铕 (Eu) 在一定条件下发生氧化还原，如：Ce 和 Eu 通常以 +3 价存在，随湿度、酸度增加， Ce^{3+} 被氧化成 Ce^{4+} ，并发生水解沉淀；在厌氧环境中， Eu^{3+} 被还原成 Eu^{2+} ，具有更高的溶解度。

稀土元素在风化、迁移或沉积过程中，易与土壤中氧化物、黏土矿物、天然有机质和矿物有机质复合物发生相互作用 (Johannesson et al.,

2004; Davranche et al., 2015)，从而产生分馏现象。基于稀土元素标准化参照数据，可采用轻稀土元素和重稀土元素含量比值、Eu 异常系数、Ce 异常系数等参数量化表征土壤稀土元素分馏特点，示踪土壤中稀土元素的地球化学行为。

2.2.1 稀土元素分馏特征

稀土元素一般可分轻稀土元素 (LREEs, La、Ce、Pr、Nd、Pm、Sm、Eu) 和重稀土元素 (HREEs, Gd、Tb、Dy、Ho、Er、Tm、Yb、Lu)。由于 Y 与重稀土元素具有相似地球化学性质，也常被划分为重稀土元素 (Li et al., 2017)。LREEs 通常被富含黏质和无定形铁氧化物的土壤和沉积物吸附，而 HREEs 易与土壤中碳酸盐结合产生沉淀 (Compton et al., 2003)。土壤溶液中 HREEs 可与有机配体发生络合作用 (Johannesson et al., 2000)，有利于稀土元素在土壤中的淋溶，促进稀土元素在土壤中富集。与 HREEs 相比，LREEs 在酸性土壤中优先淋溶，具有更高的迁移率 (Liu et al., 2019)。林地、水田、荒地、建筑用地显示多数 $(\text{La/Yb})_N$ 比值在 0.35—0.96，呈现 HREEs 更为富集状态 (Zhou et al., 2020a)。但也有研究指出农田土壤表现出 LREEs 富集状态，这种现象与农田土壤铁 / 铝氧化物优先吸附 LREEs 有关 (Tyler, 2004)。不同土地利用类型的轻重稀土元素分馏存在差异，LREEs 与森林、湿地呈显著正相关；HREEs 与农

业用地土壤显著正相关, 与城市和工业用地没有显著相关性 (Borowiak et al., 2018)。

2.2.2 Ce 异常特征

在碱性、氧化条件下, 土壤中 Ce^{3+} 易氧化为 CeO_2 (Ce^{4+}), CeO_2 不易迁移, 导致 Ce 正异常。铁锰氧化物和腐殖酸等有机物质对 Ce^{3+} 的吸附氧化能力更强 (Janots et al., 2015), 使 Ce 优先在土壤积累。铁锰氧化物对 Ce 异常的影响机制也存在争议。Cao et al. (2016) 认为 Ce 异常取决于残余原生矿物、次生黏土以及铁氧化物的性质和比例。但 Chang et al. (2016) 认为锰氧化物与 Ce 异常值呈显著负相关, 铁氧化物对 Ce 异常没有显著相关性, Ce^{3+} 氧化主要是与锰氧化物的反应引起。不同质地土壤的氧化还原条件不同, Ce 异常特征存在差异 (Li et al., 2017)。在不同土地使用类型方面, 相关研究发现仅一个林地剖面样品 Ce 负异常 (0.55—0.78) 比较明显, 水田、荒地、建筑用地等土壤剖面 Ce 和 Eu 几乎没有异常 (Zhou et al., 2020a)。

2.2.3 Eu 异常特征

Eu 具有 +3 和 +2 价, 比其他稀土元素更具流动性。与相邻元素 Sm 和 Gd 相比, 容易出现 Eu 亏损。还原条件可促进 Eu^{2+} 产生和迁移。花岗岩的风化剖面表现出 Eu 正异常, Eu^{2+} 的离子半径与 Ca^{2+} 非常相似, Eu 正异常是由蚀变过程中斜长石中的 Ca^{2+} 被 Eu^{2+} 替代引起 (Wang and Liang, 2015)。相关研究证实了石灰岩土壤中的 Eu 负异常和长英质岩石 (富含长石) 土壤中的 Eu 正异常 (Cao et al., 2016)。富长石岩石形成的土壤显示 Eu 正异常, 而云母土壤可以显示 Eu 负异常 (Laveuf and Cornu, 2009; Temga et al., 2021)。Eu 正异常也可能与碳酸盐或流体 - 矿物相互作用有关, Eu^{2+} 替代了 Ca^{2+} 、 Sr^{2+} 和 Na^+ , 导致长石中 Eu 系统性正异常 (Compton et al., 2003)。矿区周边农田土壤 Eu 异常值均值 (0.92 ± 0.10 , $n=6$) 高于对照土壤 (0.81 ± 0.04 , $n=3$) (Liu et al., 2019)。珠三角地区土壤也观察到强烈的 Eu 负异常 (Chang et al., 2016), 主要归因于还原条件下 Eu^{3+} 向 Eu^{2+} 还原, 并可能与所处的热带暖湿气候环境下易发生的强淋溶过程有关 (Laveuf and Cornu, 2009)。由此可知, 土壤 Eu 异常主要受控于两种机制: 氧化还原反应和母岩 Eu 特征。

3 土壤稀土元素主要赋存相态及其迁移富集的影响因素

3.1 土壤稀土元素主要赋存相态

土壤中稀土元素的赋存相态影响其含量特征、迁移行为和生物有效性等 (Mittermüller et al., 2016)。在土壤环境中, 稀土元素可与无机含氧酸阴离子 (CO_3^{2-} 、 SO_4^{2-} 和 PO_4^{3-} 等) 和有机配体 (腐殖酸等) 形成络合物, 主要以无机结合态、有机结合态和铁锰氧化物结合态的赋存相态分布在土壤中, 导致土壤稀土元素含量和迁移速率存在显著差异 (Feitosa et al., 2020)。

3.1.1 无机结合态

稀土离子与无机配体形成络合物与配体浓度水平、pH 值有关。硫酸盐络合态、碳酸盐络合态是土壤中稀土离子的主要赋存相态。在酸性条件下, SO_4^{2-} 是稀土元素的主要络合离子, SO_4^{2-} 与稀土离子可形成 $\text{REE}(\text{SO}_4)^+$ 、 $\text{REE}(\text{SO}_4)_2^-$ 、 $\text{REE}(\text{SO}_4)_3^{3-}$ (Gimeno Serrano et al., 2000)。当 SO_4^{2-} 浓度 $< 0.05 \text{ mol} \cdot \text{L}^{-1}$ 时, 稀土离子与 SO_4^{2-} 络合形成 $\text{REE}(\text{SO}_4)^+$, 当 SO_4^{2-} 处于饱和状态时, 稀土离子与 SO_4^{2-} 络合形成较稳定的 $\text{REE}(\text{SO}_4)_2^-$ 和稳定的 $\text{REE}(\text{SO}_4)_3^{3-}$ 。在碳酸盐浓度高的碱性环境中, 稀土离子与 CO_3^{2-} 络合作用形成 $\text{REE}(\text{CO}_3)^+$ 和 $\text{REE}(\text{CO}_3)_2^-$ 等碳酸盐络合物 (Tang and Johannesson, 2006), 也会引起轻重稀土分馏。在江西赣南稀土矿区的污染农田中, 碳酸盐结合态稀土占比高达 51% (孙峰等, 2013), 明显高于其他结合态。此外, 磷酸盐对稀土元素具有沉淀和络合作用 (Byrne et al., 1991), 土壤中添加磷酸盐可以降低稀土元素的迁移能力。

3.1.2 有机结合态

土壤中天然有机质, 如低分子量有机酸、铁载体和腐殖物质, 来源于微生物代谢产物、植物分泌物和有机质降解, 影响稀土元素迁移和生物活性 (Marsac et al., 2013)。土壤有机质含有羧基、酚羟基、N—和 S—结合位点, 易与稀土离子结合, 增强了有机质对稀土元素的络合 (Wu et al., 2001; Pourret et al., 2010)。这些络合物在氧化条件下溶解, 并释放 REEs (Davranche et al., 2015), 提高稀土元素迁移率。Li and Liu (2020) 选择铁载体和腐殖酸两种有机质, 研究了不同有机质在含稀土元素的铁氢化物溶解过程中对稀土元素迁移和分配的作用机制, 发现铁载体能催化形成

Ce 正异常，同时腐殖酸和铁载体均与 HREEs 具有高度亲和力，促进轻重稀土分馏。不同土地利用类型土壤中，有机质含量与稻田剖面和建筑场地剖面稀土元素含量具有正相关关系 (Zhou et al., 2020a)。Chang et al. (2019) 发现土壤剖面上部稀土元素容易流失，明显贫化，而下部可能由于有机质大量分解，稀土元素通过与有机胶体络合，导致稀土元素相对富集。

3.1.3 铁锰氧化物结合态

稀土元素的原子轨道壳层中具有半填充的 4f 电子结构，一般以 REE^{3+} 和 REE(OH)^{2+} 的形式吸附在铁锰氧化物表面 (Laveuf and Cornu, 2009)。铁锰氧化物对稀土离子的吸附机制为专性吸附，由于重稀土水解能力强于轻稀土，故重稀土更容易在矿物表面富集。王学锋等 (2019) 指出在江西赣南矿区周边土壤低 pH 值条件下，稀土元素容易发生水解，并与铁锰氧化物相结合，导致铁锰氧化物结合态富集。Temga et al. (2021) 报道了在铁锰氧化物富集的土壤中，稀土元素大量富集 ($6\text{--}1942 \text{ mg}\cdot\text{kg}^{-1}$)，与土壤铁和锰含量显著正相关。Mihajlovic and Rinklebe (2018) 指出铁对稀土元素含量的影响存在“门槛”值，当土壤铁含量低于 $500 \text{ mol}\cdot\text{kg}^{-1}$ 时稀土元素含量随铁含量的增加而升高；当铁含量高于 $500 \text{ mol}\cdot\text{kg}^{-1}$ 时稀土元素含量则呈现降低趋势。不同结合态的稀土元素也存在相互影响，比如：有机质的存在会影响稀土元素在黏土矿物和铁锰氧化物上的吸附 (Davranche et al., 2008)，在溶解态有机质和锰比例较低、pH 值为 3—6 时，腐殖酸与氧化锰络合稀土的竞争最为激烈 (Yang et al., 2019)。

3.2 土壤稀土元素迁移富集的影响因素

土壤中稀土元素迁移富集及分馏行为是一系列复杂的化学作用过程，包括风化淋滤、氧化还原、溶解沉淀、吸附解吸等。该过程受到土壤 pH 值、氧化还原条件和土壤质地等因素的影响和制约，导致稀土元素迁移富集以及分馏行为随着地球化学过程发生相应变化，其含量分布模式和环境迁移潜力存在差异。

3.2.1 土壤 pH 值

土壤 pH 值通过控制水 - 土反应（吸附、氧化还原反应和共沉淀等）影响稀土元素含量。pH 值不仅影响土壤中有机物和矿物的表面电性，还影响稀土离子的电离程度和形态 (Alshameri et al.,

2019)。土壤 pH 值与林地和荒地剖面的 REE 富集因子正相关，土壤 pH 对剖面稀土元素迁移具有重要影响 (Zhou et al., 2020a)。使用石灰可提高土壤 pH 值，导致更多稀土元素被带负电荷的有机物表面吸附 (Sukitprapanon et al., 2019)。Yang et al. (2019) 发现稀土在高岭石和埃洛石上的吸附量均随 pH 值增加而增加。

土壤 pH 值还影响稀土元素的分馏。稀土元素的吸附与土壤 pH 值呈正相关，随着土壤 pH 值增加，会有更多的氢氧化物离子 (OH^-) 与稀土离子形成络合物 (Cao et al., 2001)。例如：La、Ce 在酸性和还原性条件下，在土壤中更容易释放。pH 值高于 6 时，稀土元素主要以络合形式存在，且 HREEs 比 LREEs 更易被有机物络合，而 LREEs 更易于在土壤胶体中吸附和迁移。在高 pH 值和碳酸盐浓度环境下，会导致 LREEs 优先保留在土壤中，HREEs 碳酸盐复合物更易于释放和迁移 (Laveuf and Cornu, 2009)。由于 Y 和 Ho 具有相同的价态和相似的离子半径，表现出一致的地球化学行为，Y/Ho 比值一般变化不大，常用作表征稀土元素分馏的指标 (Lawrence et al., 2006)。在我国喀斯特地区，土壤 pH 值与 Y/Ho 比值呈明显的负相关关系 (Chang et al., 2019)。高 pH 值条件下，Ho 与 HCO_3^- 或有机质产生络合作用，但 Y 与氢氧化铁共沉淀作用更强，导致 Y/Ho 比率降低。在土壤剖面中，由于上部 pH 值较低阻碍了 Ho 和 HCO_3^- 络合，Ho 更容易迁移 (Quinn et al., 2006)，含 Y 相（如铁氧化物）的溶解能力增强 (Thompson et al., 2013)，可能导致剖面上部 Y/Ho 比率增加。由此可知，不同稀土元素在不同 pH 值条件下参与络合反应能力存在差异，引起稀土元素出现分馏现象。

3.2.2 氧化还原条件

氧化还原条件是控制稀土元素迁移富集行为的重要因素。通过氧化还原电位 (Eh) 可表示氧化还原能力的强弱 (王玉洁等, 2021)。土壤环境干湿更替变化、微生物活动、植物根系代谢以及外来物质的氧化还原性等会导致土壤 Eh 值改变。

在还原条件下，La、Ce、Gd 和 Y 释放增强 (Cao et al., 2001)；Eh 和 pH 值变化会引起稀土元素赋存相态发生变化，稀土元素可交换结合态和铁锰氧化物结合态浓度随 Eh 和 pH 值降低而降低。氧化还原条件通过 pH 值变化对稀土元素

释放动力学有一定的影响, 可能是由于氧化过程中质子的形成 (Cao et al., 2001; Frohne et al., 2011), Eh 增加引起 pH 值降低, 容易形成氧化和酸性双重条件, 促进稀土与铁、锰、铝的迁移和释放 (Mihajlovic et al., 2017)。在还原/氧化条件交替变化的湿地土壤环境, Fe 均为稀土元素形成的控制因素 (Guénet et al., 2018)。稀土元素在氧化阶段主要与有机质结合, 然后在还原阶段随着溶解有机物 (DOM) 的解吸而释放, 释放的有机物多数为腐殖酸、富里酸、乙酸等物质 (Fedotov et al., 2019), 使有机质成为湿地中稀土元素迁移的影响因素。

氧化还原反应过程对 Ce 异常和 Eu 异常现象具有重要解释力 (Nakada et al., 2013)。通过配体驱动的氧化还原位移会引起 Ce 正异常 (孙峰等, 2013); 而在 Eh 或 pH 值太低的情况下未能检测到 Ce 异常 (Akagi et al., 2002)。湿地植物根际代谢活动导致周边土壤环境 Eh 短期内降低, 有利于 Eu^{3+} 还原为 Eu^{2+} 而出现 Eu 负异常 (Krzciuk and Gauszka, 2020)。

3.2.3 土壤质地

土壤中稀土含量主要由母岩决定。在半干旱草原区, 母岩风化后, LREEs 和 HREEs 基本上不会发生分馏 (Huang et al., 2019)。与黄土和碳酸盐岩形成的土壤相比, 由碱性/酸性火成岩形成的土壤稀土元素含量更高 (Hu et al., 2006)。母岩成分对于区分土壤中稀土含量至关重要 (Mihajlovic et al., 2019)。Ce 异常和 Eu 异常现象也来自于对母岩特性的继承。土壤稀土元素含量随着粒径的减小和黏土粒级百分比的增加而增加, 但粒度分布只能部分解释 REEs 含量变化, 黏土含量约为 50% 时 REEs 含量相对较低, 黏土含量高于 40% 并未导致土壤剖面中稀土元素含量较高 (Compton et al., 2003)。黏土矿物对稀土元素的固定和分馏不仅取决于土壤黏土粒级含量, 还取决于土壤黏土粒级组成 (Laveuf and Cornu, 2009)。土壤稀土元素在沿地表径流迁移过程中, 与细颗粒一起长距离迁移, 而粗颗粒主要留在原位 (Zhou et al., 2020b)。Liu et al. (2019) 在研究受稀土开采活动影响的河流时发现, 在上游高 SO_4^{2-} 含量和低 pH 值条件下, 稀土元素以粒径小于 0.22 μm 的组分为主, 稀土元素形态以 REE^{3+} 、 $\text{REE}(\text{SO}_4)^+$ 和 $\text{REE}(\text{CO}_3)^+$ 为主, 随着下游区域有机质含量增加

和 pH 值升高, 稀土元素含量和形态主要由 0.22—0.45 μm 的颗粒组分决定。也有研究表明, 土壤稀土元素与黏土粒级之间没有相关性 (Feitosa et al., 2020)。

4 土壤稀土元素的生态效应

4.1 稀土元素对土壤理化性质的影响

高浓度稀土元素可引起土壤理化性质以及肥力指标变化。稀土元素输入土壤系统后, 吸附在土壤胶体表面的氢离子及铝离子会被稀土离子置换进入土壤溶液, 同时, 氢离子也可从稀土元素与有机组分络合的过程中释放出来, 增加土壤活性酸度 (Du et al., 2009)。Liang et al. (2021) 报道了稀土矿区开采后的 pH 值 (4.37—4.75) 和有机碳含量均较低, 铁和铝浓度水平高 (最高值分别为 $33 \text{ g} \cdot \text{kg}^{-1}$ 和 $103 \text{ g} \cdot \text{kg}^{-1}$)。

施加外源稀土影响土壤中氮化物的迁移与转化。鲁鹏等 (1999) 发现施加低浓度稀土微肥会促进土壤脲酶活性, 增加土壤有效氮浓度, 而施加高浓度稀土微肥则明显抑制土壤脲酶活性, 减少土壤有效氮浓度。添加外源稀土引起的土壤 pH 变化是促进有机态氮向铵态氮转化的关键因素。旱培条件下过量稀土元素很可能抑制硝化作用导致铵态氮累积和硝态氮含量降低, 淹水条件下由于土壤稀土元素抑制土壤矿化和氨化过程, 而引起氨态氮和水解氮含量减少 (丁士明等, 2004)。Wang and Liang (2014) 通过实验研究发现, 添加低浓度稀土元素时, 由于稀土离子与磷竞争吸附位点, 导致土壤溶液中磷吸附速率降低, 添加高浓度稀土元素时, 可能因为稀土离子与磷酸盐阴离子之间发生强烈的反应, 土壤中磷吸附效率提高。土壤中钾在一般以无机矿物态形式存在, 很少以有机态形式存在, 稀土元素对土壤有效钾无显著影响 (朱建国等, 2001)。

4.2 稀土元素对土壤动物生长、发育和繁殖的影响

大型动物群的多样性, 尤其是甲壳虫科和皮翅目动物的多样性, 与土壤稀土元素水平密切相关。Li et al. (2010) 研究发现土壤大型动物群落 (包括甲虫) 丰度与稀土元素水平之间呈负相关。Naccarato et al. (2020) 研究得出相反的结论, 即甲虫暴露于添加稀土元素的土壤时出现显著的 Ce、La、Y 和 Nd 积累。部分研究围绕不同稀土元素化合物对蚯蚓繁殖存活的影响机制进行了探

索。Lahive et al. (2014) 报道了蚯蚓暴露于硝酸铈铵时, 其繁殖和存活均受到负面影响, 而暴露于 CeO₂ 纳米颗粒没有出现明显影响。将红正蚓暴露于 CeO₂ 纳米颗粒物改良土壤 ($5000 \text{ mg} \cdot \text{kg}^{-1}$) 7 d 后, 其组织和粪便出现 Ce 积累 ($5.3 \text{ mg} \cdot \text{kg}^{-1}$ 和 $495 \text{ mg} \cdot \text{kg}^{-1}$) (Antisari et al., 2012)。Adeel et al. (2021) 指出稀土氧化物生物累积量呈剂量依赖性; 在 $100 \text{ mg} \cdot \text{kg}^{-1}$ 时, La₂O₃ 和 Yb₂O₃ 分别导致蚯蚓出现死亡以及繁殖减少; 超微结构观察显示, $500\text{--}1000 \text{ mg} \cdot \text{kg}^{-1}$ 的稀土氧化物可诱导内部细胞器(线粒体、高尔基体和氯粒体)异常。

4.3 稀土元素对植物生长的影响

稀土元素在土壤中积累并经常被植物吸收, 通过食物链对人体和环境健康构成潜在风险。以我国南方花岗岩红壤侵蚀流域水稻为研究对象, 其土壤稀土元素向水稻植株转移能力呈根>叶>稻谷的规律(马倩怡等, 2018)。不同施肥方式下, 喷施 CeO₂ 后大豆植株地上部分 Ce 含量大幅度增加, 而土壤施用 CeO₂ 后大豆植株地上部分 Ce 含量较低, 这种显著差异可能是受土壤中吸附-解吸、迁移过程以及根系吸收能力的影响, 导致生物有效性不同 (Salehi et al., 2018)。稀土浓度为 $0.5 \text{ mg} \cdot \text{L}^{-1}$ 和 $1 \text{ mg} \cdot \text{L}^{-1}$ 时, 根部长度、株高和生物量显著降低 (Martinez et al., 2018)。从植物细胞角度, 稀土元素可定位在植物细胞的细胞壁、质膜、细胞间隙、中柱导管部位, 一定剂量稀土元素还能引起细胞核和内质网形态学变化, 使其具有非正常形态; 高浓度 La³⁺ 破坏叶绿体超微结构, 降低叶绿体矿质元素和叶绿素含量, 抑制植物光合作用和生长 (Hu et al., 2016); CeO₂ 纳米颗粒悬浮液可改变超氧化物歧化酶 (SOD) 活性, 诱导脂质过氧化和细胞膜损伤, 抑制根系生长 (Cui et al., 2014)。在外源因素影响下, 添加不同铁盐可诱导 LREEs、中稀土元素 (MREEs) 和 HREEs 产生不同毒性效应, 添加氯化亚铁 (Ⅱ) 时 LREEs 对水稻植株生长起抑制作用, 添加硫酸亚铁 (Ⅱ) 时 MREEs 和 HREEs 对水稻生长也具有毒性效应 (Martinez et al., 2018); 稀土元素在酸雨、氨氮、盐、重金属和干旱等胁迫下对种子萌发、植物生长、产量和质量以及生理生化过程也具有影响 (Liu et al., 2016; Zhang et al., 2017; Liu et al., 2021a), 其影响程度取决于胁迫因子强度和稀土元素浓度组合水平。

4.4 稀土元素对微生物的影响

高含量稀土元素对土壤微生物种群结构和多样性的影响显著, 从而影响土壤生态系统稳定性。主要有三种表现: 一是稀土元素对细菌抑制作用都很强 (罗建美和季宏兵, 2005); 二是长期暴露在原地淋滤采矿矿区土壤环境, 可能会迫使微生物种群进化为具有更高耐受性和抗性的群落 (Liu et al., 2021b); 三是稀土元素污染会抑制不具有耐性的真菌生长, 降低真菌群落结构多样性, 具有耐性的菌株由于其适应性强而大量繁殖生长 (袁浩等, 2019)。同时, 添加稀土元素影响土壤微生物群落代谢水平, 导致土壤 pH 变化。稀土作用于土壤微生物群落和酶, 间接产生生物反应, 导致土壤氮素形态和含量变化 (丁士明等, 2004)。稀土尾矿土壤中营养物质变化, 特别是硫酸盐水平的增加和有效氮水平的降低, 导致了真菌群落组成和多样性较低, 以及以真菌为主的异养硝化作用发生改变 (Li et al., 2019)。离子型稀土矿区经过地浸开采后, 细菌、古菌和氨氧化古菌丰富度与稀土元素呈显著负相关, 稀土元素和铵是影响细菌群落结构的主导因子 (Liu et al., 2021b)。不同矿区环境异质性较高, 稀土复合污染可引起细菌群落丰富度和多样性、细菌群落组成发生明显变化。

4.5 土壤中稀土元素潜在健康风险

通过环境暴露及食物链传递, 稀土元素在人体的累积产生较大的健康风险。一方面是呼吸暴露, 土壤与大气环境之间易于产生稀土元素颗粒物迁移转化, 通过实验设计大鼠吸入 CeO₂ 纳米颗粒后发现, 增加了肺泡巨噬细胞产生的白细胞介素 -12 和干扰素 -γ (Ma et al., 2011; Ma et al., 2014)。稀土矿区人群通过吸入 PM_{2.5} 摄入的稀土元素平均剂量为 $5.09 \times 10^{-7}\text{--}2.25 \times 10^{-5} \text{ mg} \cdot \text{kg}^{-1} \cdot \text{d}^{-1}$, 儿童从道路粉尘中摄入的 LREEs 平均剂量为 $4.27 \times 10^{-4}\text{--}2.63 \times 10^{-2} \text{ mg} \cdot \text{kg}^{-1} \cdot \text{d}^{-1}$, 比通过其他暴露途径 (如吸入 PM_{2.5}) 高 2—3 个数量级, 且受粒径大小影响显著, 儿童从道路灰尘的细颗粒 ($<100 \mu\text{m}$) 中摄取 LREEs 具有严重风险 (Li et al., 2020)。另一方面是食物链传递, 在福建稀土矿区附近和江西地区的蔬菜样品中, 稀土元素含量分别为 $3.58 \mu\text{g} \cdot \text{g}^{-1}$ (干重) (Li et al., 2013) 和 $6.37 \mu\text{g} \cdot \text{g}^{-1}$ (干重) (Zhuang et al., 2017)。广州稀土元素

高丰度地区蔬菜中稀土元素均值达 $24.95 \mu\text{g} \cdot \text{g}^{-1}$ (干重) (Wang et al., 2022), 土壤稀土元素向植物迁移富集后通过食物链被人体摄入。马倩怡等 (2018) 指出稻谷中稀土元素含量虽未超过日均稀土元素摄入量限值 ($5.33 \mu\text{g} \cdot \text{kg}^{-1} \cdot \text{d}^{-1}$), 但人体长期食用仍存在一定的健康风险。长期低剂量摄入稀土元素, 对脑、肝、骨、内分泌和免疫功能等具有非常广泛的生物学毒效应, 能导致多种毒害和病变 (Rim et al., 2013; Aalapati et al., 2014)。

4.5.1 分子和细胞水平潜在健康风险

在分子水平, 镧系元素与血细胞/生物分子的生物化学相互作用可以在吸收、转运及其组织摄取机制方面发挥重要作用 (Ansoborlo et al., 2006)。对人体和动物的研究表明, Ce 改变了血红素的氧亲和力 (Cheng et al., 2000)。由于 Ce^{4+} 的电荷-离子半径比最接近 Fe^{3+} , Ce^{4+} 与血红素相互作用能够较大程度地影响血红素的氧结合作用 (Kumar et al., 2016)。稀土元素对体

外细胞影响的研究结果如表 2 所示。在细胞水平, 细胞凋亡已被认为是污染物对细胞毒性效应的重要表现 (Heller et al., 2019), 不同稀土元素对同一细胞系的敏感性及细胞毒作用存在显著差异 (Feyerabend et al., 2010)。Heller et al. (2019) 通过体外实验发现肾细胞系对镧系元素 (La、Ce、Eu 和 Yb) 具有浓度和时间依赖性效应, 体外环境微量浓度下镧系元素对肾细胞没有显著影响, 但强调了在采矿和工业区高浓度 Ce 具有潜在有害影响。Wu et al. (2020) 观察到 LaCl_3 对体外培养的 bEnd.3 细胞具有杀伤作用, 并伴有细胞内 Ca^{2+} 的增加。Korotkov et al. (2020) 发现 Gd^{3+} 可抑制线粒体内膜通透性转换孔 (MPTP) 的开放。线粒体是细胞凋亡的主要执行细胞器, 在凋亡发生过程中, 随着线粒体功能和形态改变, 导致细胞内活性氧 (ROS) 水平升高, 进而通过DNA损伤的方式诱导细胞凋亡 (Noorimotlagh et al., 2018)。

表 2 稀土元素对体外细胞的影响
Tab. 2 Effect of rare earth elements on cells *in vitro*

细胞系列 Cell series	目标 Object	结果 Result	参考文献 References
HL-60 细胞、NB4 细胞 HL-60 cells, NB4 cells	研究 LaCl_3 和 CeCl_3 对细胞生长和凋亡影响机制 To study the mechanism of LaCl_3 and CeCl_3 on cell growth and apoptosis	在一定浓度下, 稀土化合物抑制白血病细胞生长, 诱导其凋亡, 对正常骨髓造血祖细胞 (CFU-GM) 没有明显的抑制作用。 At a certain concentration, rare earth compounds inhibit the growth and induce apoptosis of leukemia cells, and have no obvious inhibitory effect on normal bone marrow hematopoietic progenitor cells (CFU-GM).	Dai et al., 2002
外周血单个核细胞 Peripheral blood mononuclear cells	研究稀土元素对酶活性和细胞凋亡的影响 To study the effects of rare earth elements on enzyme activity and apoptosis	暴露于低浓度稀土, 外周血单个核细胞端粒酶活性高于对照组, 对细胞凋亡率无影响。 Under low concentration rare earth exposure, the telomerase activity of peripheral blood mononuclear cells was higher than that of the control group, and had no effect on the apoptosis rate.	Yu et al., 2007
HeLa 细胞 HeLa cells	研究柠檬酸镧 (LaCit) 对细胞诱导的失巢凋亡影响 To study the effect of Lanthanum Citrate (LaCit) on cell induced anoikis apoptosis	出现稀土元素诱导细胞失巢凋亡, 显示出剂量依赖性细胞反应。 Rare earth elements induced cell loss of nest and apoptosis, showing a dose-dependent cellular response.	Su et al., 2009
人骨肉瘤 (MG63)、人脐带血管周围 (HUCPV) 细胞、小鼠巨噬细胞 (RAW 267.4) Human osteosarcoma (MG63), human umbilical cord perivascular (HUCPV) cells, mouse macrophages (RAW 267.4)	研究暴露于 Y、La、Ce、Pr、Nd、Eu、Gd 后的细胞活性 To study the cell activity after exposure to Y, La, Ce, Pr, Nd, Eu and Gd	细胞活力下降, 胞浆中单核和寡核小体富集, 不同镧系元素对同一细胞系的敏感性及细胞毒作用存在显著差异。 Cell viability decreased and mononuclear and oligonucleosomes were enriched in the cytoplasm. There were significant differences in the sensitivity and cytotoxicity of different lanthanides to the same cell line.	Feyerabend et al., 2010

(待续 To be continued)

(续表 2 Continued Tab. 2)

细胞系列 Cell series	目标 Object	结果 Result	参考文献 References
MC3T3-E1 细胞 MC3T3-E1 cells	研究暴露于 CeCl ₃ 溶液后 MC3T3-E1 细胞的碱性磷酸酶 (ALP) 活性 To study the alkaline phosphatase (ALP) activity of MC3T3-E1 cells after exposure to CeCl ₃ solution	CeCl ₃ 对成纤维细胞有刺激作用, 但对成骨细胞有抑制作用。加入 rhBMP-2 能够使 ALP 活性高于对照水平, rhBMP-2 可促进矿物沉积, 控制 Ce ³⁺ 浓度水平以降低毒性。 CeCl ₃ can stimulate fibroblasts, but inhibit osteoblasts. Adding rhBMP-2 can make ALP activity higher than the control level. rhBMP-2 can promote mineral deposition and control Ce ³⁺ concentration level to reduce toxicity.	Schmidlin et al., 2012
HeLa 细胞 HeLa cells	评估低浓度 GdCl ₃ 对细胞增殖的影响机制 To evaluate the mechanism of low concentration GdCl ₃ on cell proliferation	Gd 通过诱导 HeLa 细胞在浓度小于 100 μmol·L ⁻¹ 时进入 S 期而促进增殖。 Gd induces HeLa cells at concentrations less than 100 μmol·L ⁻¹ enters S phase and promotes proliferation.	Zhang et al., 2009
U937 人组织细胞、人外周血单核细胞 (PBMCs) U937 human tissue cells, human peripheral blood monocytes (PBMCs)	研究镧系元素的细胞毒性 To study the cytotoxicity of lanthanides	镧系元素仅在非常高浓度下具有细胞毒性, 在较低浓度下细胞具有良好耐受性。 Lanthanides are cytotoxic only at very high concentrations and well tolerated at lower concentrations.	Bladen et al., 2013
Jurkat 细胞、人外周血淋巴细胞 (HPL) Jurkat cells, human peripheral blood lymphocytes (HPL)	评估镧 (La) 对两种细胞系的细胞毒性和遗传毒性 To evaluate the cytotoxicity and genotoxicity of lanthanum (La) on two cell lines	镧对两种细胞系均具有细胞毒性和遗传毒性, HPL 对 La 处理比 Jurkat 细胞更敏感, 坏死是 La 诱导细胞毒性的途径。氧化应激可能参与了基因毒性过程。 Lanthanum has cytotoxicity and genotoxicity to both cell lines. HPL is more sensitive to La treatment than Jurkat cells. Necrosis is the pathway of La induced cytotoxicity. Oxidative stress may be involved in the process of genotoxicity.	Paiva et al., 2009
肾大鼠 NRK-52E、人 HEK-293 细胞 Renal rat NRK-52E cells, human HEK-293 cells	研究 La、Ce、Eu 和 Yb 暴露对两种哺乳动物肾细胞系活力的影响 To study the effects of La, Ce, Eu and Yb exposure on the viability of two mammalian kidney cell lines	镧系元素对两种细胞系均显示浓度和时间依赖性影响, 其中 Ce 是最有效的元素。在体外环境痕量浓度下对肾细胞影响不明显。 Lanthanides showed concentration and time-dependent effects on both cell lines, of which Ce was the most effective element. The concentration of trace renal cells had no significant effect <i>in vitro</i> .	Heller et al., 2019

4.5.2 个体水平潜在健康风险

在个体水平, 主要通过小鼠模拟实验总结相应结论。La、Ce 和 Nd 显著抑制小鼠肾脏中抗氧化酶基因和蛋白质的表达, 肌酐增加, 尿酸、尿素氮、钙和磷减少等表征肾功能被破坏的现象, 表明小鼠暴露于 REEs 后发生肾炎或上皮细胞坏死与氧化应激密切相关; 肾脏损伤程度按 Ce³⁺、Nd³⁺、La³⁺ 暴露依次降低, 这可能归因于镧系元素 4f 电子结构差异 (Zhao et al., 2013)。Cheng et al. (2014) 对小鼠灌胃给药 CeCl₃ 持续 90 d, 小鼠白细胞和中性粒细胞减少, 碱性磷酸酶、乳酸脱氢酶、胆碱酯酶、甘油三酯和胆固醇增加, 在涉及多个端点的 675 个基因中观察到差异表达。稀土元素对特定生长时期的大鼠健康状况也会产生影响, 比如高剂量硝酸钇诱导怀孕 Sprague-Dawley 大鼠及其后代的外周血氧化应激和细胞凋亡, 同时抗氧化能力增强 (Zhang et al., 2020); 大鼠从胚胎期至断奶后 1 个月期间给予不同浓度

氯化镧 (LaCl₃), 显示 La 可损害大鼠空间学习记忆, 促进线粒体分裂, 导致海马神经细胞线粒体分裂融合障碍, La 还能诱导过多的有丝分裂吞噬, 提示 La 引起的学习记忆障碍可能与 MQC 紊乱有关 (Yu et al., 2020)。

5 未来研究展望

由于土壤环境污染错综复杂以及稀土元素来源多样化, 目前对土壤环境中稀土元素行为及风险的认识尚未深入清晰。稀土元素的分馏和赋存相态显著影响了稀土元素在土壤中的迁移过程及生态环境效应, 因而也影响着稀土元素在土壤环境中产生的生物生态风险和人体健康风险。未来需进一步探讨和研究以下问题:

(1) 根据不同时空尺度跟踪调查土壤稀土元素的来源与贡献, 深入研究不同时间跨度稀土元素在土壤 - 水 - 沉积物 - 作物系统中的相关性, 从大尺度空间差异的视角揭示土壤稀土元素积累、

分配和长距离迁移规律, 为土壤中稀土元素污染防治和资源高效利用提供科学依据。

(2) 部分研究已经关注到土壤稀土元素与重金属以及不同稀土元素之间的复合污染问题, 需进一步研究土壤稀土元素与其他新兴污染物交互作用和协同污染迁移机制, 评估稀土元素复合污染的剂量-生物响应机制及健康风险, 为土壤稀土元素环境持久性、生物累积性和毒性等风险评估提供支撑依据。

(3) 探究微生物对稀土元素富集环境的响应策略, 加强土壤微生物对稀土元素积累、分馏和迁移规律的影响研究, 解析微生物介导铁矿还原及二次成矿过程对稀土元素赋存相态和富集迁移的影响及机制, 可为稀土矿资源化利用和稀土污染土壤生物修复提供理论支撑。

(4) 研制土壤稀土元素高效清洁的吸附材料, 从源头控制土壤稀土元素迁移和污染, 提高稀土元素的开采利用效率; 强化土壤稀土元素连续监测以及对开采盗采稀土矿行为的监督管理, 从稀土开采、运输、加工、资源化利用全过程中管控稀土元素去向; 研发稀土元素利用新技术, 净化回收土壤富集的稀土元素。

(5) 未来的研究应进一步阐明全球范围不同类型土壤(高动态湿地等)稀土元素释放动力学的控制因素以及潜在的活化过程, 对于揭示稀土元素的生物地球化学循环的区域异质性规律和全球化过程均具有重要启示。

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